

青海南山当家寺花岗岩体锆石 U-Pb 年代学、 地球化学及其地质意义

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内容提要:青海南山构造带发育北西—南东向展布的印支期花岗岩带, 其研究对厘清古特提斯演化阶段西秦岭造山带与南祁连构造带的衔接转换关系具有重要意义。当家寺花岗岩体是由花岗闪长岩和二长花岗岩组成的复合花岗岩基。本文对当家寺花岗岩体进行了详细的岩石学、岩石地球化学和 LA-ICP-MS 锆石 U-Pb 同位素年代学研究。结果表明, 花岗闪长岩和二长花岗岩的结晶年龄分别为 240.1 ± 2.1 Ma 和 241.0 ± 2.6 Ma, 属于中三叠世。地球化学研究显示岩体相对高硅($\text{SiO}_2 = 66.37\% \sim 73.99\%$)和富钾($\text{K}_2\text{O} = 3.37\% \sim 4.73\%$), 属于准铝质—弱过铝质高钾钙碱性 I 型花岗岩。岩石稀土配分曲线表现为轻稀土富集的右倾型, 具中等程度的铕负异常($\delta\text{Eu} = 0.32 \sim 0.64$)。微量元素显示富集大离子亲石元素 Cs、Rb、K, 相对亏损高场强元素 Nb、Ta、P、Ti。岩石成因研究表明其形成于下地壳基性岩为主的源岩的部分熔融, 同时存在一定程度幔源岩浆贡献。结合区域地质背景资料, 认为当家寺花岗岩体形成于洋壳俯冲阶段, 可能为宗务隆洋向南俯冲的地球动力学背景下的产物。

关键词:青海南山; 西秦岭北缘; 当家寺花岗岩体; 锆石 U-Pb 定年; 地球化学; 构造环境

“中央造山系”是呈东西向横亘于中国大陆中央的巨型造山系, 其东段为东秦岭-大别造山带; 西段分为西秦岭造山带、祁连造山带和昆仑造山带, 中间夹有中祁连、欧龙布鲁克(全吉)和柴达木等微陆块(Zhang Guowei et al., 1998, 2004; Lu Songnian et al., 2002; Lu Songnian et al., 2006; Xu Zhiqin et al., 2006)。中国中西部地区造山带发育的多条蛇绿(混杂)岩、岛弧火山岩和相应的变质基底及沉积盖层展现了从原特提斯洋至古特提斯洋演化阶段“多微陆块/多洋(海)盆/多岛弧”的复杂构造格局, 构造演化表现为多个微陆块之间的裂离与聚合(Pan Guitang et al., 1997; Yin Hongfu et al., 1998; Lu Songnian et al., 2004; Shi Rendeng et al., 2004; Xu Zhiqin et al., 2006; Wang Zongqi et al.,

2009; Chen Yibing et al., 2010; Liu Chengjun et al., 2014; Zhang Zhaowei et al., 2015; Sun Jiaopeng et al., 2015; Li Zhao et al., 2016; Li Ruibao et al., 2016; Yang Zhangzhang et al., 2016; Dai Xiong, et al., 2016; Guo Xianqing et al., 2016)。因此中国中西部诸造山带及夹于其中的微陆块的衔接、拼合关系成为特提斯构造域研究的关键科学问题。

西秦岭造山带西部的共和盆地作为西秦岭-松潘大陆构造结中次一级的西南构造结点, 处于西秦岭、东昆仑、祁连、柴达木及欧龙布鲁克等多个构造带及块体交接转换的重要结点地区(Zhang Guowei et al., 2004)。共和盆地为一新生代盆地(Chang Hong et al., 2009)目前对于其构造归属有坳拉谷(Sun Yangui et al., 2004; Zhang Guowei et al.,

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2004)、西秦岭楔(Yan Zhen et al., 2012)和西秦岭造山带西延(Feng Yimin et al., 2003)等不同的认识。同时对共和盆地周缘广泛发育的印支期花岗岩的成因也存在争议,一是岩体形成的构造体制是阿尼玛卿-勉略洋向北俯冲碰撞的结果(Guo et al., 2012; Luo Biji et al., 2012; Luo et al., 2012; Wei Ping et al., 2013; Jin Xiaoye et al., 2013; Huang Xiongfei et al., 2014),或是宗务隆洋俯冲的产物(Guo Anlin et al., 2009; Peng Yuan et al., 2016),亦或是以先期存在于柴达木地块与欧龙布鲁克微地块之间的早古生代蛇绿岩带为基础发生的陆壳俯冲(Zhang Hongfei et al., 2006);二是岩体形成的构造环境是俯冲环境或俯冲板片的断离(Zhang Hongfei et al., 2006; Guo Anlin et al., 2009; Wei Ping et al., 2013; Peng Yuan et al., 2016)还是后碰撞阶段的地壳加厚或岩石圈拆沉(Jin Weijun et al., 2005; Zhang Hongfei et al., 2006; Zhang Chengli et al., 2008; Yin Yong et al., 2009; Li Ting et al., 2012; Xu Xueyi et al., 2012, 2014; Sun Xiaopan et al., 2013; Zhang Tao et al., 2014; Yang Shuanhai et al., 2015);目前对于西秦岭地区印支晚期处于后碰撞阶段的认识分歧较小(Zhang Hongfei et al., 2006; Zhang Chengli et al., 2008; Liu Zhipeng et al., 2012; Li Zuochen et al., 2013; Ren Houzhou et al., 2014; Zhang Tao et al., 2015; Yang Yang et al., 2015; Lu Dongyu et al., 2015),但 Huang Xiongfei et al. (2014)对印支早期西秦岭地区存在加厚地壳背景下的埃达克岩提出质疑。因此,进一步开展共和盆地周缘花岗岩的研究,对于揭示西秦岭造山带印支期造山过程,厘清古特提斯演化阶段中央造山系西段各造山带及块体的衔接转换关系等具有重要理论意义。

本文选取青海南山构造带东段龙羊峡水库北侧的当家寺花岗岩体为研究对象,通过野外详细观察、系统采样及室内测试分析,提供了该岩体的全岩地球化学和精确的锆石 U-Pb 测年结果,结合区域上前人地质研究成果,对该岩体岩石成因及动力学背景进行讨论,试图为认识西秦岭北缘、宗务隆构造带及南祁连构造带交接部位的印支期构造格局和深部地球动力学过程提供证据。

1 区域地质概况及岩体地质特征

青海南山构造带是位于共和盆地北缘,呈北西—南东向展布,其北侧以青海湖南山断裂为界与

南祁连构造带相接,西端与柴达木盆地北缘地区宗务隆构造带相接,为衔接西秦岭造山带、南祁连构造带和宗务隆构造带的交接部位。青海南山构造带内出露的主要地层包括属于南祁连构造带化隆微地块的古元古代化隆岩群中深变质岩系,该变质岩系主体为黑云石英片岩、二云石英片岩,夹二云母斜长片麻岩和黑云斜长片麻岩;属于宗务隆构造带的石炭纪—二叠纪中浅变质岩系,岩石组合为黑云石英片岩夹含石榴石黑云石英片岩及黑云母变粒岩、大理岩、长石石英岩等;属于西秦岭北缘构造带的下三叠统隆务河组,该套地层由砾岩、含砾砂岩、砂岩、泥岩及灰岩共同组成,发育正粒序层理、平行层理、波纹层理、包卷层理以及底模构造等浊流沉积等典型构造标志,同时可见滑塌褶曲。在龙羊峡一带,该套沉积组合总体由砂岩、粉砂岩、黑色泥岩和薄层灰岩局部夹透镜状砾屑灰岩或砾岩构成,具有浊流和碎屑流沉积特征;新近系、第四系在区内分布广泛。区内侵入岩主要为印支早期基性、中性及中酸性侵入岩,属于西秦岭北缘印支期岩浆岩带的一部分。基性侵入岩以辉长岩、辉石岩为主,多与同期的中酸性侵入岩伴生,区内主要的基性侵入岩体包括与黑马河花岗岩体伴生的辉长岩体以及江西沟地区的拉日陇哇、拉日托陇和拉木陇哇等辉长岩体;中性侵入岩以闪长岩、石英闪长岩为主,多分布于区内主要的花岗岩基内;中酸性侵入岩以花岗闪长岩和二长花岗岩为主,研究区出露有黑马河、江西沟和当家寺等大型花岗岩基。依据野外接触关系和同位素年代学资料,总体上基性侵入岩的形成时代早于中性侵入岩,中性侵入岩的形成时代早于中酸性侵入岩。区内侵入岩多侵位于下三叠统隆务河组中。

当家寺花岗岩体位于共和盆地的东北缘,青海南山构造带东段(图 1a),岩体整体略呈椭圆状,长轴方向为 NNW—SSE,出露面积约 290km²,其主体侵位于下三叠统隆务河组(T₁L),但其周围多被新近系及第四系覆盖。岩体主要由花岗闪长岩和二长花岗岩组成,其中二长花岗岩分布于岩体中部,花岗闪长岩分布于岩体边部,二者呈渐变接触关系(图 1b),野外未见明显接触界线。接触带两侧的花岗闪长岩和二长花岗岩多具似斑状结构,花岗闪长岩中斑晶以钾长石为主,二长花岗岩中斑晶以斜长石为主,花岗闪长岩和二长花岗岩均具有中粒—中粗粒结构,内部无明显相变。花岗闪长岩和二长花岗岩中均发育少量暗色微粒包体(图 2g),但含量和规模均较小,包体岩性多为闪长质,直径一般 3~5

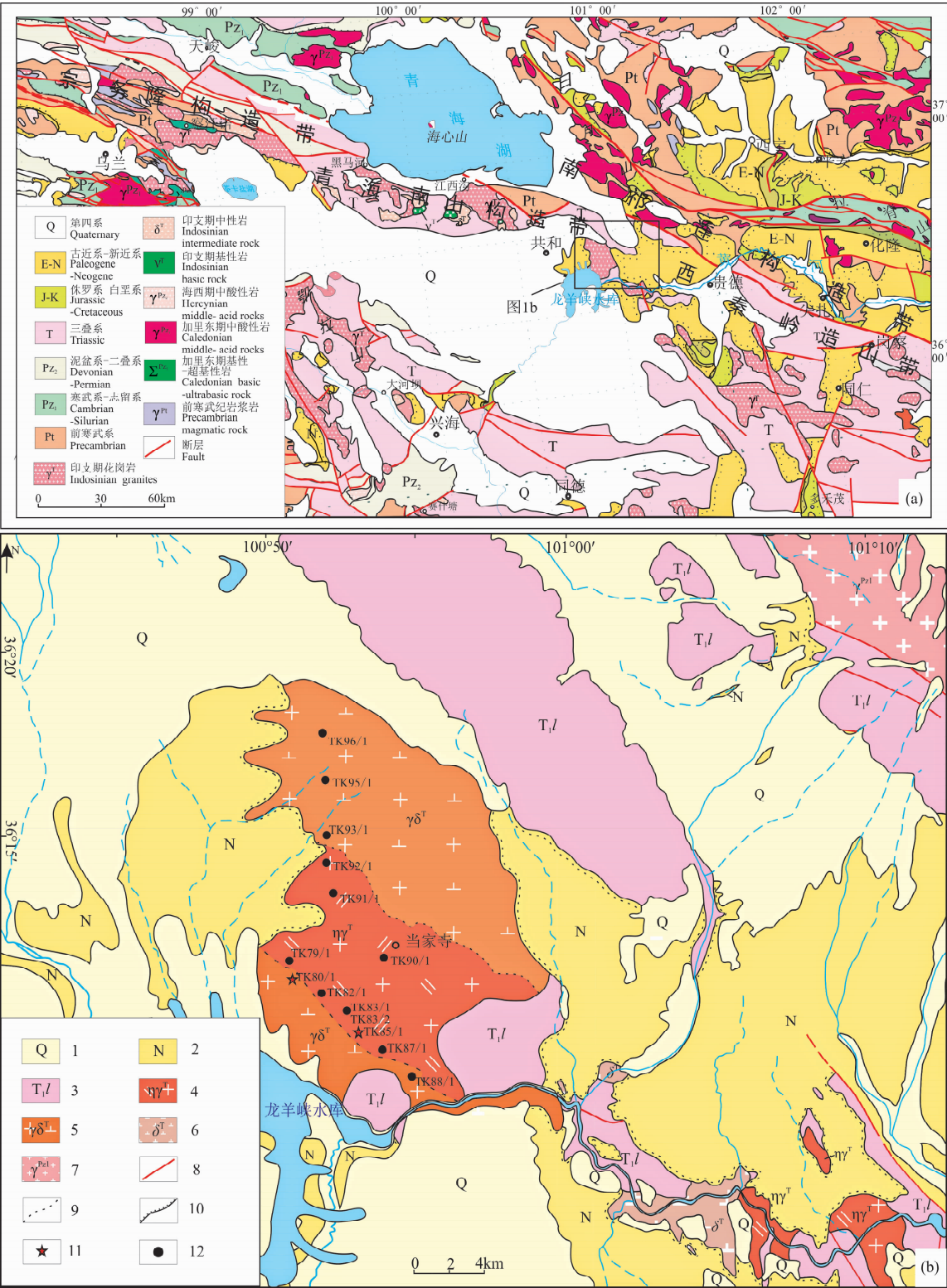


图 1 青海南山当家寺花岗岩体构造位置(a)及岩体地质简图(b)

Fig.1 Regional location map (a) and geological map (b) of Dangjiashi granitic complex in the Qinghai Nanshan tectonic zone
1—第四系;2—新近系;3—下三叠统隆务河组;4—三叠纪二长花岗岩;5—三叠纪花岗闪长岩;6—三叠纪闪长岩;7 早古生代花岗岩;
8—断层;9—岩相界线;10—角度不整合界线;11—同位素采样;12—全岩采样
1—Quaternary;2—Neogene;3—Lower Triassic Longwuhe Formation;4—Triassic monzonite granite;5—Triassic granodiorite;
6—Triassic diorite;7—Early Paleozoic granite;8—fault;9—boundaries of lithofacies;10—angular unconformity;
11— isotopes sample position;12—sample position

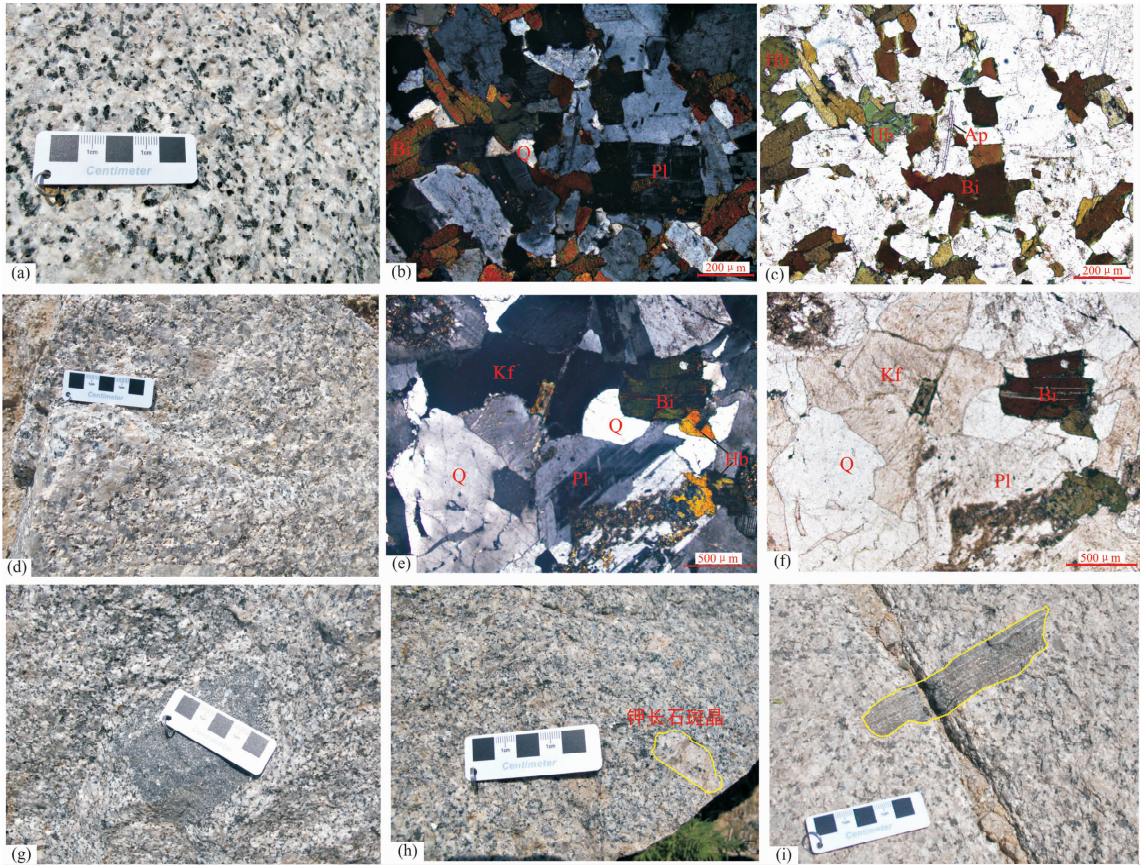


图 2 当家寺花岗岩体野外露头及显微镜下照片

Fig. 2 Filed and photomicrographs of representative samples of the Dangjiasi granitic complex

(a)、(b)、(c)—花岗闪长岩；(d)、(e)、(f)—二长花岗岩；(g)—花岗闪长岩中包体；(h)—似斑状二长花岗岩中钾长石斑晶；

(i)—二长花岗岩中的围岩捕掳体；Pl—斜长石；Q—石英；Kf—钾长石；Bi—黑云母；Hb—角闪石；Ap—磷灰石

(a)、(b)、(c)—granodiorite；(d)、(e)、(f)—monzonite granite；(g)—enclaves in granodiorite；

(h)—crystals of feldspar in porphyritic monzonite granite；(i)—xenolith in monzonite granite；Pl—plagioclase；

Q—quartz；Kf—K-feldspar；Bi—biotite；Hb—hornblende；Ap—apatite

cm,小者约 1 cm,大者可达 20 cm。包体形态多样,多呈不规则状、水滴状或椭圆状,与寄主岩界线截然。岩体中偶见围岩捕掳体(图 2i)。

花岗闪长岩呈灰白色,中粒花岗结构,块状构造,主要由斜长石(50%~55%)、钾长石(10%~15%)、石英(20%±)、角闪石(5%±)和黑云母(5%~10%)组成;副矿物有磁铁矿、锆石、磷灰石、榍石、褐帘石等。斜长石呈半自形板状,大小一般 3~5 mm,部分 2~3 mm,局部被绢云母及少量方解石交代;钾长石为半自形板状,偶见卡氏双晶,粒径以 2~5 mm 为主;石英呈他形粒状,填隙状分布,大小一般 2~5 mm,部分 5~8 mm;角闪石为浅绿色半自形、他形柱状,大小一般 0.2~1 mm;黑云母棕褐色—棕黄色多色性,呈片状。具似斑状结构的花岗闪长岩,斑晶主要为钾长石(1%~5%),呈半自形板状,大小一般 5~10 mm。基质呈中粒花岗结构,主

要由斜长石、钾长石、石英、黑云母组成,矿物粒径一般为 2~5 mm。

二长花岗岩呈浅肉红色,中粗粒花岗结构,块状构造,主要矿物为钾长石(35%~40%)、斜长石(30%~35%)、石英(25%~30%)、黑云母(5%±);副矿物包括磁铁矿、锆石、磷灰石、榍石等。钾长石呈半自形宽板状,粒度大小一般 2~5 mm,部分 5~8 mm,发育卡式双晶,少数颗粒表面高岭土化;斜长石呈半自形板状,部分颗粒可见环带结构,颗粒大小以 3~5 mm 为主,部分 5~6 mm;石英颗粒大小在 2~5 mm 之间,呈半自形、他形粒状,填隙状分布;黑云母呈自形、半自形片状,片直径一般 1~2 mm,镜下具有黄褐色—棕褐色多色性,少数黑云母已风化蚀变为绿泥石。具似斑状结构的二长花岗岩的斑晶为石英(5%~10%)和斜长石(5%±),均为半自形晶,大小一般 5~10 mm,基质主要由斜长石、钾长石、石英、角闪石、

黑云母组成,大小一般 2~3 mm。

2 样品采集与分析方法

2.1 样品采集

为了保证样品的代表性及研究的科学性,本次野外工作选择了由南至北穿越岩体路线,沿线系统采集了花岗闪长岩和二长花岗岩的地球化学样品和锆石 U-Pb 定年样品。其中,锆石 U-Pb 定年样品 2 件,地球化学样品 14 件。同位素年龄采样点地理坐标为:花岗闪长岩(TK80/1)36°10′57.8″N,100°51′48.7″E;二长花岗岩(TK85/1)36°11′21.8″N,100°51′38.3″E(图 1b)。

2.2 锆石 U-Pb 定年

用于年代学研究的样品由河北省廊坊市峰泽源岩矿检测技术有限公司完成粉碎和锆石的分离工作。锆石制靶及阴极发光照相工作由北京锆年领航科技有限公司完成。锆石 U-Pb 测年在天津地质矿产研究所通过 Neptune 质谱仪利用 LA-ICP-MS 方法完成测定,首先根据锆石阴极发光照片、反射光和透射光照片选择锆石的合适的测年晶域,再利用 193 nm 激光器对锆石进行剥蚀,通常采用的激光剥蚀的斑束直径为 35 μm ,以 TEMORA 作为外部锆石年龄标准。采用中国地质大学刘勇胜博士研发的 ICPMS Data Cal 程序和 Kenneth R. Ludwig 的 Isoplot 程序进行数据处理,利用²⁰⁸Pb 校正法对普通铅进行校正。以 NIST612 玻璃标样作为外标计算锆石样品的 Pb、U、Th 含量。详细分析方法及仪器参数见 Li Huaikun et al. (2009)。

2.3 岩石地球化学分析

全岩主量、稀土和微量元素测试在长安大学西部矿产资源与地质工程教育部重点实验室完成。主量元素测试采用 XRF 法,测定流程包括:①烧失量计算:将坩锅在烘箱内 150℃干燥 3 h 后,称其重量 W_1 ,加入约 1 g 样品,称样品重量 W_2 ;然后放入 900℃的马弗炉中 8 h,降温后放入干燥器静置 20 min,随后称重得 W_3 。通过公式 $(\text{LOI}) = (W_1 + W_2 - W_3) / W_2$ 计算出样品的烧失量(LOI)。②玻璃熔融法制样:主量元素测定时首先称取样品 0.50 g,以无水四硼酸锂和硝酸铵为氧化剂,倒入铂金坩锅中,再加入适量溴化锂,在 1200℃左右振荡熔融制成玻璃薄片。③使用 X 射线荧光光谱仪测定。稀土和微量元素分析采用 Thermo-X7 电感耦合等离子体质谱仪,分析精度和准确度优于 10%。将 200 目以下样品(500 mg)置于 PTFE 坩锅,加入添加剂(1.0

mL 高纯 HF 和 1.5 mL 高纯 HNO₃),按照标准测试程序,反复添加、加热、冷却后,最后在离心管中稀释到 50 mL;将所得溶液在电感耦合等离子体质谱仪(ICP-MS)上完成测定。

3 分析结果

3.1 锆石 U-Pb 测年

花岗闪长岩样品(TK80/1)中锆石多数晶粒为无色透明至淡黄色,呈自形短柱状或长柱状,一般长约 100~200 μm ,宽约 60~110 μm 。阴极发光图像显示大多数锆石具有典型的岩浆韵律环带(图 3a);锆石的 U 含量为 $398 \times 10^{-6} \sim 522 \times 10^{-6}$,Th 含量为 $123 \times 10^{-6} \sim 290 \times 10^{-6}$,Th/U 的比值为 0.20~0.71(表 1),显示其具有岩浆锆石特征(Corfu et al., 2003)。所有测点都尽量选择锆石边部的震荡环带区,在 24 个测点中,有 2 个测点(7、22 号测点)可能由于锆石封闭体系遭破坏,使放射成因铅丢失而明显远离谐和线,因而不参与计算。其余 22 分析测点的²⁰⁶Pb/²³⁸U 和²⁰⁷Pb/²³⁵U 谐和性较好,²⁰⁶Pb/²³⁸U 年龄值介于 232~251 Ma 之间,²⁰⁶Pb/²³⁸U 加权平均年龄值为 240.1 ± 2.1 Ma(MSWD=0.79)(图 4a),代表了花岗闪长岩的结晶年龄。

二长花岗岩样品(TK85/1)中锆石多为自形程度较高的长柱状晶体,少数为短柱状,长轴约 100~310 μm ,短轴约 60~130 μm 。阴极发光图像显示清晰的震荡环带(图 3b),属于典型的岩浆成因锆石。锆石的 U 含量为 $380 \times 10^{-6} \sim 531 \times 10^{-6}$,Th 含量为 $213 \times 10^{-6} \sim 348 \times 10^{-6}$,Th/U 的比值为 0.40~0.92,均大于 0.1(表 1),表明属于岩浆锆石(Corfu et al., 2003)。本次测试了 24 个测点,由于放射性 Pb 丢失造成 3 个测点(11、12、24 号测点)远离谐和线(Duncan et al., 1999),剔除 3 个测点后,其余 21 个测点均投影于谐和线上或其附近,²⁰⁶Pb/²³⁸U 年龄值变化于 235~250 Ma 之间,加权平均年龄为 241.0 ± 2.6 Ma(MSWD=1.2)(图 4b),代表了岩石结晶年龄。因此,当家寺花岗岩体侵位年龄为 240~241 Ma,为中三叠世。

3.2 主量元素特征

花岗闪长岩 SiO₂ 含量为 66.37%~69.86%,样品相对富钾,K₂O 为 3.37%~4.29%,平均为 3.82%,K₂O/Na₂O(0.79~1.59),平均为 1.3;Na₂O+K₂O 为 6.4%~8%,碱度率 AR=1.96~3.12;岩石 Al₂O₃、TF₂O₃、MgO 和 TiO₂ 含量相对较高,Al₂O₃ 为 14.15%~15.84%,TF₂O₃ 为

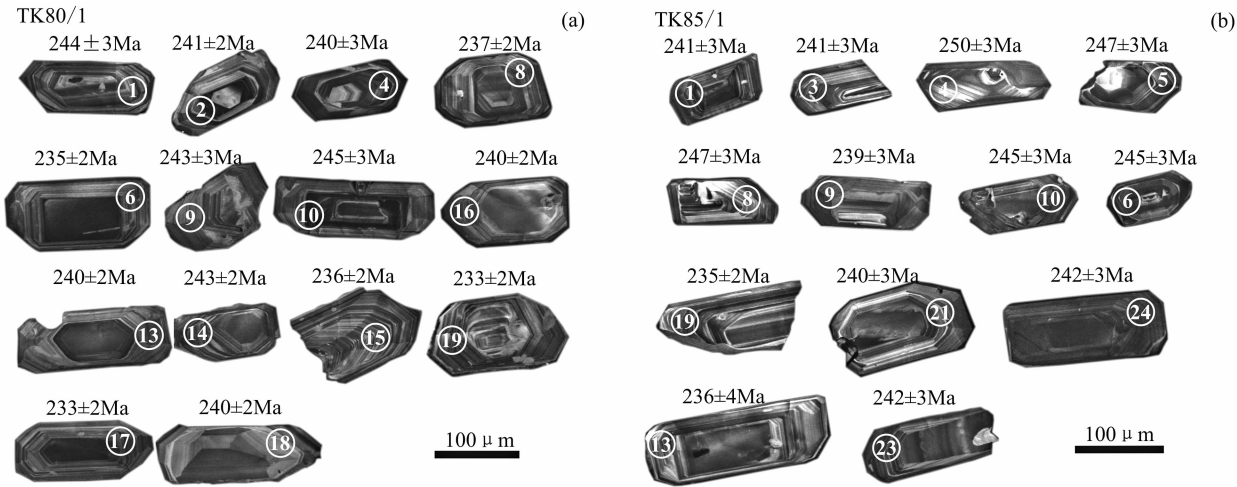


图 3 当家寺花岗岩体花岗闪长岩(TK80/1)(a)和二长花岗岩(TK85/1)(b)代表性单颗粒锆石阴极发光图像
Fig. 3 CL images of selected zircons for granodiorite (TK80/1) (a) and monzonitic granite (TK85/1) (b) of the Dangjiasi granitic complex

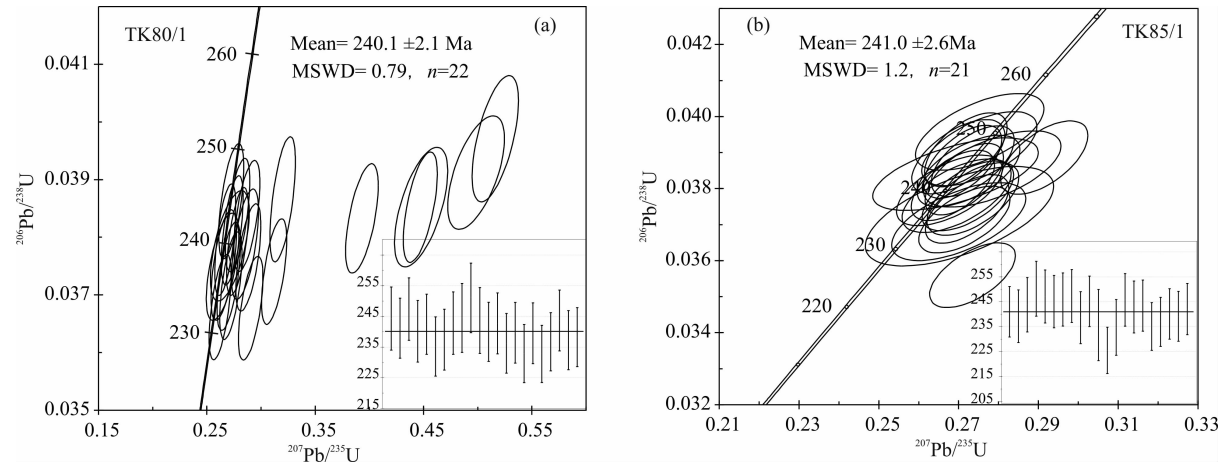


图 4 当家寺花岗岩体花岗闪长岩(TK80/1)(a)和二长花岗岩(TK85/1)(b)锆石 U-Pb 年龄谐和图
Fig. 4 U-Pb concordia diagrams of zircon for granodiorite (TK80/1) (a) and monzonitic granite (TK85/1) (b) of the Dangjiasi granitic complex

3.09%~4.11%,MgO 为 1.16%~1.39%,TiO₂ 为 0.39%~0.51%,平均为 0.45%(表 2)。在 QAP 岩石分类图解上(图 5a)样品主体落入花岗闪长岩区。样品里特曼指数介于 1.67~2.40 之间,为钙碱性系列,与 AFM 图解(图 5b)及 SiO₂-K₂O 图解(图 6a)所反映的信息一致,所有样品均落在高钾钙碱性区域;岩体 A/CNK 值介于 1.00~1.11 之间,大部分小于 1.10,平均为 1.04,表现为弱过铝质特征(图 6b)。分异指数 DI 介于 72.22~88.95 之间。

二长花岗岩 SiO₂ 含量相对较高,为 69.58%~73.99%,Al₂O₃ 变化范围较大(12.92%~14.89%);样品富钾特征明显,Na₂O+K₂O(6.7%~8%),K₂O/Na₂O 均值为 1.3,碱度率 AR=2.43

~3.12;TF₂O₃、MgO、TiO₂ 和 CaO 含量相对花岗闪长岩为低,TF₂O₃ 为 1.69%~3.2%,MgO 为 0.42%~0.91%,CaO 为 1.44%~2.80%,TiO₂ 为 0.19%~0.37%(表 2),在 QAP 岩石分类图解上(图 5a)都落入二长花岗岩范围内。样品里特曼指数为 1.53~2.46,在 AFM 图解(图 5b)及 SiO₂-K₂O 图解(图 6a)上落入高钾钙碱性区域;岩体 A/CNK 值介于 0.89~1.06 之间,平均为 1.04,具有准铝质—弱过铝质特征(图 6b)。

当家寺花岗岩体在哈克图解(图 7)上,均表现为 TiO₂、Al₂O₃、TF₂O₃、MgO、CaO、P₂O₅ 随着 SiO₂ 含量的增加而降低的趋势,Na₂O 的变化趋势不明显,K₂O 含量随着 SiO₂ 含量增加而增加。

表 1 当家寺花岗岩体花岗闪长岩 (TK80/1) 和二长花岗岩 (TK85/1) LA-ICP-MS 锆石 U-Pb 同位素分析结果

Table 1 LA-ICP-MS zircon U-Pb analytic data for Dangjiasi granitic complex

测点	含量($\times 10^{-6}$)		Th/U	同位素比值						年龄(Ma)					
	Th	U		$^{206}\text{Pb}/^{238}\text{U}$	1 σ	$^{207}\text{Pb}/^{235}\text{U}$	1 σ	$^{207}\text{Pb}/^{206}\text{Pb}$	1 σ	$^{206}\text{Pb}/^{238}\text{U}$	1 σ	$^{207}\text{Pb}/^{235}\text{U}$	1 σ	$^{207}\text{Pb}/^{206}\text{Pb}$	1 σ
TK80/1(花岗闪长岩)															
1	211	522	0.40	0.0386	0.0002	0.2725	0.0032	0.0513	0.0006	244	1	245	3	254	23
2	245	485	0.50	0.0381	0.0001	0.2735	0.0023	0.0520	0.0004	241	1	245	2	287	14
3	236	443	0.53	0.0391	0.0001	0.4985	0.0087	0.0932	0.0018	247	1	411	6	1492	37
4	251	481	0.52	0.0380	0.0001	0.2670	0.0021	0.0510	0.0004	240	1	240	2	243	12
5	290	409	0.71	0.0383	0.0001	0.3932	0.0036	0.0745	0.0007	242	1	337	3	1054	19
6	221	514	0.43	0.0372	0.0001	0.2722	0.0024	0.0531	0.0004	235	1	244	2	345	17
7	200	467	0.43	0.0405	0.0001	0.5647	0.0047	0.1012	0.0007	256	1	455	3	1647	14
8	243	487	0.50	0.0375	0.0001	0.2655	0.0034	0.0514	0.0006	237	1	239	3	257	30
9	219	513	0.43	0.0384	0.0001	0.2784	0.0030	0.0526	0.0005	243	1	249	2	322	22
10	123	610	0.20	0.0387	0.0002	0.3200	0.0026	0.0604	0.0006	245	1	282	2	617	20
11	275	398	0.69	0.0397	0.0002	0.5162	0.0058	0.0939	0.0007	251	1	423	4	1507	15
12	267	420	0.64	0.0385	0.0002	0.4477	0.0083	0.0836	0.0013	244	1	376	6	1283	34
13	231	501	0.46	0.0379	0.0001	0.2783	0.0022	0.0532	0.0004	240	1	249	2	339	23
14	224	506	0.44	0.0384	0.0001	0.2869	0.0033	0.0542	0.0006	243	1	256	3	389	19
15	246	489	0.50	0.0373	0.0001	0.2702	0.0028	0.0525	0.0005	236	1	243	2	309	22
16	258	475	0.54	0.0379	0.0001	0.2852	0.0036	0.0546	0.0007	240	1	255	3	398	32
17	233	508	0.46	0.0368	0.0001	0.2623	0.0025	0.0517	0.0005	233	1	237	2	272	22
18	249	488	0.51	0.0379	0.0001	0.2712	0.0065	0.0519	0.0012	240	1	244	5	283	58
19	259	478	0.54	0.0368	0.0001	0.2904	0.0028	0.0573	0.0005	233	0	259	2	502	20
20	245	483	0.51	0.0374	0.0001	0.3122	0.0026	0.0606	0.0005	237	1	276	2	633	19
21	283	437	0.65	0.0385	0.0001	0.4470	0.0031	0.0844	0.0007	244	1	375	2	1302	17
22	237	428	0.55	0.0386	0.0001	0.5863	0.0049	0.1101	0.0009	244	1	469	3	1811	15
23	254	475	0.54	0.0375	0.0001	0.2645	0.0051	0.0512	0.0010	237	1	238	4	256	44
24	223	504	0.44	0.0377	0.0001	0.2883	0.0030	0.0555	0.0006	238	1	257	2	435	22
TK85/1(二长花岗岩)															
1	348	380	0.92	0.0381	0.0001	0.2686	0.0026	0.0512	0.0005	241	3	242	2	250	19
2	275	463	0.59	0.0378	0.0002	0.2695	0.0044	0.0517	0.0008	239	3	242	4	272	37
3	271	459	0.59	0.0385	0.0002	0.2784	0.0061	0.0519	0.0010	244	3	249	5	280	41
4	243	496	0.49	0.0396	0.0002	0.2753	0.0056	0.0503	0.0010	250	3	247	4	209	44
5	311	418	0.74	0.0391	0.0002	0.2718	0.0025	0.0505	0.0005	247	3	244	2	220	22
6	336	391	0.86	0.0387	0.0001	0.2710	0.0021	0.0508	0.0004	245	3	244	2	232	21
7	277	458	0.60	0.0389	0.0002	0.2699	0.0033	0.0503	0.0006	246	3	243	3	209	26
8	255	472	0.54	0.0391	0.0002	0.2752	0.0048	0.0510	0.0008	247	3	247	4	239	32
9	217	523	0.41	0.0377	0.0002	0.2736	0.0032	0.0526	0.0006	239	3	246	3	322	26
10	254	481	0.53	0.0388	0.0001	0.2735	0.0025	0.0512	0.0005	245	3	246	2	250	53
11	273	425	0.64	0.0386	0.0001	0.4950	0.0060	0.0931	0.0012	244	3	408	4	1500	29
12	278	428	0.65	0.0405	0.0001	0.4665	0.0060	0.0838	0.0012	256	3	389	4	1288	28
13	213	531	0.40	0.0372	0.0004	0.2702	0.0090	0.0516	0.0011	236	4	243	7	333	48
14	300	437	0.69	0.0356	0.0001	0.2735	0.0027	0.0557	0.0005	225	2	245	2	443	20
15	284	451	0.63	0.0371	0.0002	0.2700	0.0060	0.0523	0.0007	235	3	243	5	298	1
16	278	456	0.61	0.0389	0.0001	0.2915	0.0026	0.0544	0.0004	246	3	260	2	387	49
17	253	492	0.51	0.0384	0.0001	0.2692	0.0078	0.0508	0.0015	243	3	242	6	232	67
18	286	455	0.63	0.0385	0.0001	0.2725	0.0040	0.0513	0.0007	243	3	245	3	254	27
19	336	401	0.84	0.0371	0.0001	0.2715	0.0032	0.0530	0.0006	235	2	244	3	328	26
20	311	424	0.73	0.0374	0.0001	0.2721	0.0026	0.0528	0.0005	237	2	244	2	320	20
21	273	462	0.59	0.0380	0.0001	0.2677	0.0042	0.0512	0.0008	240	3	241	3	250	35
22	252	486	0.52	0.0378	0.0001	0.2723	0.0032	0.0523	0.0006	239	3	245	3	298	26
23	267	468	0.57	0.0383	0.0001	0.2714	0.0021	0.0515	0.0004	242	3	244	2	261	10
24	218	508	0.43	0.0383	0.0001	0.3115	0.0027	0.0590	0.0005	242	3	275	2	565	17

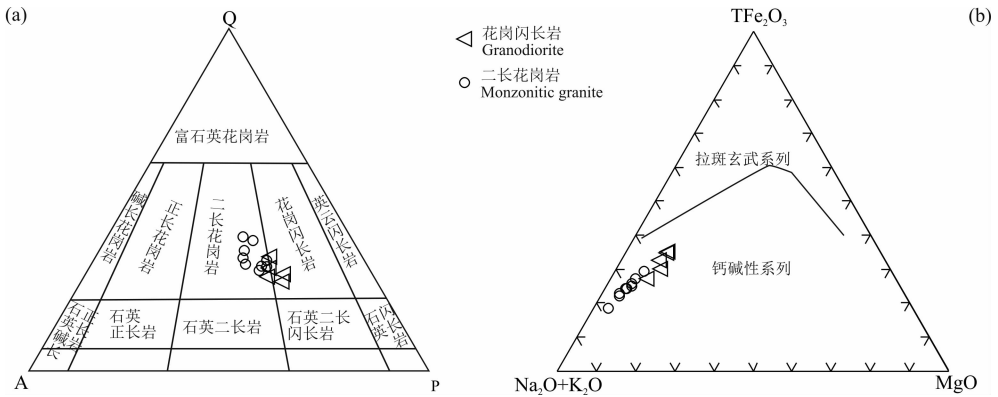


图 5 当家寺花岗岩体 QAP 图解(a)及 AFM 图解(b)(据 Irvine and Baragar,1971)

Fig. 5 QAP diagrams (a) and AFM diagrams (b) for Dangjiasi granitic complex (after Irvine and Baragar,1971)

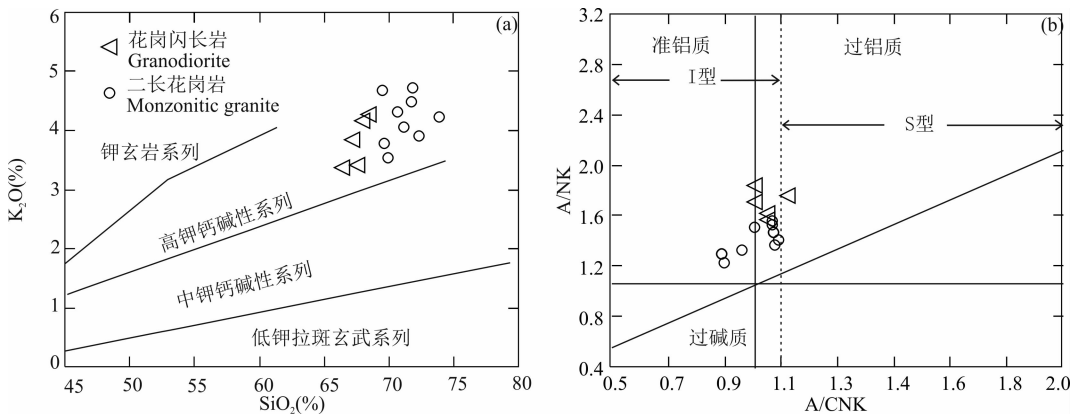


图 6 当家寺花岗岩体 SiO_2 - K_2O 图解(a,据 Rickwood,1989)和 A/NK - A/CNK 图解(b,据 Maniar and Piccoli,1989, 虚线代表 I 型和 S 型花岗岩之间的边界,据 Chappell and White,1992)

Fig. 6 SiO_2 vs. K_2O diagram (a, after Rickwood,1989) and A/NK - A/CNK diagram (b, after Maniar and Piccoli, 1989; dashed line represents boundary between I- and S-type granitoides, after Chappell and White, 1992) for Dangjiasi granitic complex

3.3 稀土与微量元素特征

由表 3 可以看出,当家寺花岗岩体稀土总量变化较大,其中花岗闪长岩稀土总量($\sum \text{REE}$)为 $175.14 \times 10^{-6} \sim 223.73 \times 10^{-6}$,平均 205.47×10^{-6} ,二长花岗岩稀土总量稍高($\sum \text{REE}$)为 $167.67 \times 10^{-6} \sim 282.39 \times 10^{-6}$,平均 228.06×10^{-6} 。球粒陨石标准化稀土配分图解中表现为轻稀土富集、重稀土亏损的右倾型(图 8), $(\text{La}/\text{Yb})_{\text{N}} = 8.31 \sim 15.97$,轻重稀土强烈分异; $(\text{La}/\text{Sm})_{\text{N}}$ 值为 $4.37 \sim 6.39$,轻稀土内部分异较明显; $(\text{Gd}/\text{Yb})_{\text{N}}$ 值为 $1.22 \sim 2.02$,中重稀土分异不明显,重稀土稀土配分曲线具有相对平坦并略微上翘的特征; $\delta \text{Eu} = 0.32 \sim 0.64$,呈现出中等程度 Eu 负异常。

微量元素原始地幔标准化蛛网图解上(图 8),岩体富集大离子亲石元素(Cs,Rb,Th,K),亏损高场强元素(Nb,Ta,Ti),具 Zr,Hf 正异常。随着分异程度

的增高(花岗闪长岩→二长花岗岩),总体表现为不相容元素逐渐富集,Sr,P,Ti 亏损程度增加,Cr,Co,Ni,V 等相容元素丰度随着 SiO_2 含量增加而降低,反映岩浆经历了一定程度的分离结晶作用。岩体 Nb/Ta 为 $7.24 \sim 23.21$;Sr/Y 比值较低,均值为 9.75;Rb/Sr 均值为 0.92;La/Y 值较低,为 $0.87 \sim 1.55$ 。

4 讨论

4.1 岩石源区与成因

当家寺花岗岩体含有角闪石以及副矿物组合中的榍石、磁铁矿;岩体具有较低的 P_2O_5 含量($0.05\% \sim 0.10\%$),并且随 SiO_2 的增加而呈现明显的降低趋势(图 7),A/CNK 小于 1,这些特征显示岩体具有 I 型花岗岩特征(Wolf and Wyllie, 1994; Chappell,1999;Wu et al.,2003;Wu Fuyuan et al., 2007)。岩体富集 Rb,Th,Cs 等大离子亲石元素

表 2 当家寺花岗岩体主量元素组成 (%)

Table 2 Major elements compositions (%) of the Dangjiassi granitic complex

样品号	花岗闪长岩					二长花岗岩								
	TK80/1	TK88/1	TK93/1	TK95/1	TK96/1	TK79/1	TK82/1	TK83/1	TK83/2	TK85/1	TK87/1	TK90/1	TK91/1	TK92/1
SiO ₂	67.17	68.45	66.37	68.00	67.55	71.78	69.58	70.75	69.86	69.56	71.19	72.37	73.99	71.88
TiO ₂	0.41	0.39	0.51	0.45	0.50	0.28	0.31	0.31	0.37	0.37	0.28	0.29	0.19	0.25
Al ₂ O ₃	14.85	15.30	15.84	15.03	15.13	14.02	14.89	14.13	14.51	14.15	13.97	13.69	12.92	14.49
TFe ₂ O ₃	3.46	3.09	4.11	3.79	4.02	2.48	2.72	2.70	3.20	3.20	2.58	2.48	1.69	2.35
MnO	0.05	0.05	0.07	0.07	0.06	0.03	0.04	0.04	0.05	0.05	0.03	0.04	0.02	0.06
MgO	1.39	1.16	1.33	1.22	1.30	0.64	0.78	0.72	0.87	0.91	0.60	0.66	0.42	0.60
CaO	2.79	2.69	3.93	3.40	2.73	1.97	2.36	2.24	2.80	2.60	2.22	2.25	1.44	1.67
Na ₂ O	3.04	3.12	3.03	2.62	3.00	2.87	2.84	3.62	4.46	3.22	4.27	2.78	2.98	3.18
K ₂ O	3.85	4.29	3.37	4.17	3.42	4.50	4.69	4.32	3.54	3.78	4.07	3.92	4.23	4.73
P ₂ O ₅	0.09	0.10	0.10	0.09	0.10	0.06	0.06	0.05	0.06	0.07	0.05	0.05	0.05	0.06
LOI	1.04	0.73	0.68	1.34	1.72	0.71	0.56	0.68	0.93	2.04	0.58	0.67	0.69	0.63
Total	98.14	99.37	99.34	100.18	99.53	99.34	98.83	99.56	100.65	99.95	99.84	99.20	98.62	99.90
K ₂ O/Na ₂ O	1.27	1.37	1.11	1.59	1.14	1.57	1.65	1.19	0.79	1.17	0.95	1.41	1.42	1.49
AR	2.28	2.40	1.96	2.17	2.12	2.71	2.55	2.88	2.72	2.44	3.12	2.45	3.02	2.92
A/NK	1.62	1.56	1.83	1.70	1.75	1.46	1.53	1.33	1.30	1.51	1.22	1.55	1.36	1.40
ACNK	1.04	1.04	1.00	1.00	1.11	1.06	1.06	0.96	0.89	1.00	0.90	1.06	1.07	1.08
σ	1.96	2.16	1.75	1.84	1.68	1.89	2.13	2.27	2.38	1.84	2.47	1.53	1.68	2.17
DI	77.65	79.47	72.22	75.61	76.75	84.89	82.22	84.81	83.09	80.9	86.91	83.44	88.95	86.45

注:TFe₂O₃是全铁含量;A/NKC=(Al₂O₃)/(CaO+K₂O+Na₂O)摩尔分数比,A/NK=(Al₂O₃)/(K₂O+Na₂O)摩尔分数比;里特曼指数σ=(K₂O+Na₂O)²/(SiO₂-43)(参见 Deng Jinfu et al. , 2015b)。

(LILE), 亏损 Nb、Ta、Ti 等高场强元素(HFSE), 具陆壳或弧岩浆特征(Thompson et al. , 1984; Zheng Yongfei et al. , 2013)。同时, 当家寺花岗岩体较低的 Sr 含量和 La/Yb、Sr/Y 比值, 不具有埃达克质岩石特征, 说明其不可能为俯冲洋壳部分熔融的产物。Ba 相对于 Rb 和 Th 亏损明显, 体现出的是成熟度较高的陆壳岩石特征(Ma Changqian et al. , 2004)。当家寺花岗岩体中花岗闪长岩 Nb/Ta 比值介于 10.95~16.68 之间, 平均 13.31; 二长花岗岩 Nb/Ta 比值介于 7.25~23.21 之间, 平均 14.69, 均值较接近大陆地壳比值(10~14)(Sun et al. , 1989; Zhao Zhenhua et al. , 2008), 暗示岩浆主要为地壳部分熔融形成。实验岩石学已证明地壳中玄武质岩石的部分熔融可以产生偏基性的准铝质花岗岩类(Beard and Lofgren, 1991; Rapp and Watson, 1995; Johannes and Holtz, 1996; Sisson et al. , 2005), 而地壳中碎屑岩类的部分熔融可以产生偏酸性的过铝质花岗岩类(Johannes and Holtz, 1996; Patino-Douce and Harris, 1998; Patino-Douce and McCarty, 1998)。当家寺花岗岩体总体呈现准铝质—弱过铝质特征, 岩体存在中等程度的负 Eu 异常, 同时具有较为平坦的重稀土元素分布模式, 可能反映的是下地壳基性岩为主的源岩发生部分熔融作用, 且源区残留矿物包含有一定数量的斜长石和角

闪石(Beard and Lofgren, 1991; Rushmer, 1991; Tepper et al. , 1993)。实验岩石学表明, CaO/Na₂O 比值介于 0.3~1.5 的花岗质岩石来源于变杂砂岩或火成岩(Jung et al. , 2007)。当家寺花岗岩体 CaO/Na₂O 比值为 0.48~1.3, 平均为 0.8; 同时在图 9a 中岩体样品主体落入角闪岩熔融区域, 由此推断当家寺花岗岩体的源区可能以变基性岩为主。岩体 Nb/Ta 比值有较宽的变化范围, 说明岩体的原始岩浆可能受到了幔源岩浆的影响。岩体中发育暗色微粒包体, 也佐证了幔源岩浆的贡献; 岩体中包体含量较少, 显示幔源岩浆对岩体成分影响有限, 幔源岩浆对岩体的影响更多的是提供热源, 可能为地幔楔受俯冲流体交代发生部分熔融形成的镁铁质岩浆底侵下地壳的结果。岩体 Al₂O₃ 含量较低, 具有明显 Eu 负异常, 显示岩体形成的深度可能较浅(Deng Jinfu et al. , 1995; Deng et al. , 2004)。当家寺花岗岩体相对高 K(Rb)(Rb 平均含量 188.34×10⁻⁶), 其原因可能类似于多数西秦岭花岗岩, 是继承了下地壳高 K 的特征(Zhang Hongfei et al. , 2005), 与高 K 的岛弧玄武质下地壳的角闪石和黑云母的脱水熔融反应有关(Sisson et al. , 2005)。

当家寺花岗岩体中花岗闪长岩、二长花岗岩具有演化的主微量元素特征, 说明两者可能为同一岩浆演化的产物。在 Harker 图解上(图 7)岩体显示

表 3 当家寺花岗岩体稀土和微量元素组成($\times 10^{-6}$)

Table 3 Trace element abundance ($\times 10^{-6}$) of the Dangjiasi granitic complex

样品号	花岗闪长岩					二长花岗岩								
	TK80/1	TK88/1	TK93/1	TK95/1	TK96/1	TK79/1	TK82/1	TK83/1	TK83/2	TK85/1	TK87/1	TK90/1	TK91/1	TK92/1
Li	108.22	114.22	68.69	73.97	121.53	94.18	77.75	86.88	105.12	88.40	82.32	58.61	96.77	143.88
Be	2.71	2.71	2.01	1.96	2.42	2.69	2.03	1.69	2.23	2.48	2.56	2.06	2.93	2.88
Sc	8.12	5.81	10.13	9.91	10.44	3.55	8.67	8.04	10.46	9.90	6.65	6.59	1.19	3.09
V	36.92	32.83	40.47	34.31	37.53	19.21	20.90	18.74	23.02	23.57	16.01	16.75	12.61	17.92
Cr	30.16	18.14	14.93	15.15	13.54	7.12	9.62	12.97	10.47	11.05	7.70	8.65	3.81	5.56
Co	7.77	6.63	8.31	7.52	8.22	5.00	4.94	4.61	5.34	5.38	3.86	4.82	2.66	3.58
Ni	6.14	10.76	12.51	13.25	12.59	13.43	12.79	13.62	13.15	15.49	14.45	14.38	15.74	15.19
Cu	11.90	1.57	2.99	1.88	1.75	3.56	2.54	3.24	9.48	12.36	2.31	9.40	1.72	0.75
Zn	42.98	47.54	49.25	45.89	48.40	24.54	36.59	40.62	49.30	39.63	42.73	38.35	21.35	38.79
Ga	16.71	17.24	17.70	15.98	17.03	14.82	16.53	15.85	17.52	16.87	16.92	15.65	13.81	15.63
Rb	200.62	200.45	149.96	172.65	155.24	190.22	204.35	160.66	159.99	180.64	221.09	171.89	217.87	251.27
Sr	241.96	316.19	281.65	264.45	307.40	188.71	185.33	179.70	195.62	293.72	150.82	172.04	152.10	173.81
Y	29.73	25.46	30.53	33.81	34.87	31.72	38.47	32.43	42.66	44.48	49.06	32.28	24.39	29.13
Zr	190.13	199.12	239.83	221.85	269.88	211.54	217.42	225.55	268.59	238.07	206.55	229.01	169.36	179.17
Nb	13.73	13.48	12.39	11.82	13.17	13.43	13.12	13.23	16.51	15.11	15.47	11.97	11.33	15.23
Mo	0.49	0.62	0.73	0.72	0.64	0.46	0.42	0.49	0.57	0.58	0.50	0.59	0.62	0.60
Cd	0.12	0.13	0.15	0.15	0.17	0.21	0.15	0.14	0.19	0.14	0.13	0.22	0.12	0.11
In	0.05	0.06	0.04	0.05	0.05	0.02	0.06	0.06	0.08	0.09	0.07	0.05	0.04	0.06
Cs	17.00	22.25	8.71	5.04	16.59	14.60	12.87	7.84	9.18	8.50	14.82	11.36	14.95	30.21
Ba	445.58	496.10	482.94	610.94	400.43	378.95	568.13	575.08	347.07	540.53	368.01	380.11	224.30	273.37
La	44.32	39.52	42.82	29.51	40.06	43.32	53.74	39.49	48.75	38.95	44.03	44.47	32.37	33.52
Ce	93.02	81.89	83.36	61.78	83.97	86.51	114.61	81.83	108.66	83.24	94.25	93.60	69.39	68.66
Pr	6.69	6.17	6.32	4.91	6.18	6.46	8.63	6.33	8.04	6.49	7.42	6.86	4.97	5.03
Nd	30.84	29.36	29.98	24.66	29.20	29.38	41.39	31.09	38.73	32.17	36.20	32.93	22.28	23.05
Sm	4.36	4.62	4.49	4.24	4.47	4.38	6.15	4.89	5.97	5.33	5.96	4.84	3.24	3.60
Eu	0.79	0.87	0.94	0.89	0.88	0.63	0.79	0.78	0.76	0.79	0.65	0.70	0.46	0.51
Gd	4.24	4.19	4.36	4.31	4.52	4.22	5.84	4.79	6.06	5.68	6.17	4.70	2.95	3.44
Tb	0.60	0.57	0.63	0.66	0.68	0.61	0.85	0.72	0.92	0.90	0.95	0.68	0.44	0.52
Dy	3.56	3.13	3.69	4.08	4.07	3.69	4.89	4.20	5.49	5.43	5.93	4.06	2.62	3.15
Ho	0.70	0.59	0.74	0.81	0.83	0.74	0.95	0.82	1.09	1.11	1.19	0.81	0.53	0.64
Er	2.16	1.74	2.19	2.37	2.51	2.27	2.80	2.35	3.13	3.29	3.54	2.39	1.65	2.07
Tm	0.31	0.25	0.32	0.35	0.37	0.35	0.37	0.32	0.43	0.48	0.51	0.34	0.26	0.32
Yb	2.10	1.67	2.08	2.40	2.46	2.40	2.53	2.04	2.75	3.25	3.35	2.16	1.84	2.27
Lu	0.32	0.25	0.31	0.36	0.38	0.39	0.38	0.31	0.42	0.47	0.51	0.32	0.29	0.35
Hf	3.93	4.05	4.18	4.13	5.12	4.47	4.45	4.31	5.34	4.71	4.44	4.45	3.69	3.92
Ta	1.22	1.23	0.74	0.86	0.94	1.37	0.77	0.57	0.86	0.97	1.05	0.70	1.56	1.82
Pb	21.18	28.47	20.14	22.49	21.60	22.11	22.71	22.64	22.00	20.69	22.35	20.46	25.65	32.91
Bi	0.08	0.22	0.05	0.04	0.04	0.35	0.09	0.04	0.20	0.09	0.04	0.07	0.02	0.11
Th	18.53	18.25	14.24	13.78	19.86	23.03	20.07	12.36	15.05	13.54	22.92	16.38	21.39	25.64
U	2.02	2.13	1.29	2.59	2.50	3.16	1.18	0.98	1.17	1.31	2.01	1.13	2.27	2.62
ΣREE	223.74	200.28	212.76	175.14	215.44	217.06	282.39	212.39	273.88	232.06	259.69	231.13	167.67	176.27
(La/Yb) _N	14.26	15.97	13.89	8.31	10.98	12.18	14.34	13.05	11.95	8.07	8.87	13.91	11.84	9.95
δEu	0.56	0.60	0.64	0.63	0.59	0.44	0.40	0.49	0.38	0.44	0.32	0.44	0.44	0.44
Nb/Ta	11.27	10.96	16.68	13.73	13.94	9.82	16.99	23.21	19.25	15.58	14.70	17.13	7.25	8.35
Nb/U	6.80	6.31	9.58	4.55	5.28	4.25	11.14	13.47	14.12	11.52	7.68	10.61	5.00	5.82
Sr/Y	8.14	12.42	9.22	7.82	8.82	5.95	4.82	5.54	4.59	6.60	3.07	5.33	6.24	5.97

注:δEu =Eu_{CN}/(Sm_{CN}×Gd_{CN})^{0.5},下标 CN 表示用球粒陨石标准化。

存在较明显的岩浆演化趋势,随着 SiO₂增加,MgO、TF_{e2}O₃、Al₂O₃、TiO₂、P₂O₅减少,可能与角闪石、斜长石、Ti-Fe 氧化物和磷灰石的分离结晶有关。二长花岗岩 Eu 负异常高于花岗闪长岩,亦说明岩浆存在着斜长石的分离结晶;这也与 Rb/Sr-Sr 图解(图 9b)所反映的信息一致(Peccerillo et al.,1976)。

二长花岗岩的 Rb/Sr 比值和稀土元素含量高于花岗闪长岩,从微量元素方面印证了二长花岗岩和花岗闪长岩为同一岩浆演化的结果。

4.2 构造环境

青海南山构造带存在广泛的印支早期岩浆活动,造成大量壳幔岩浆的侵位,形成了区内辉石岩、辉长

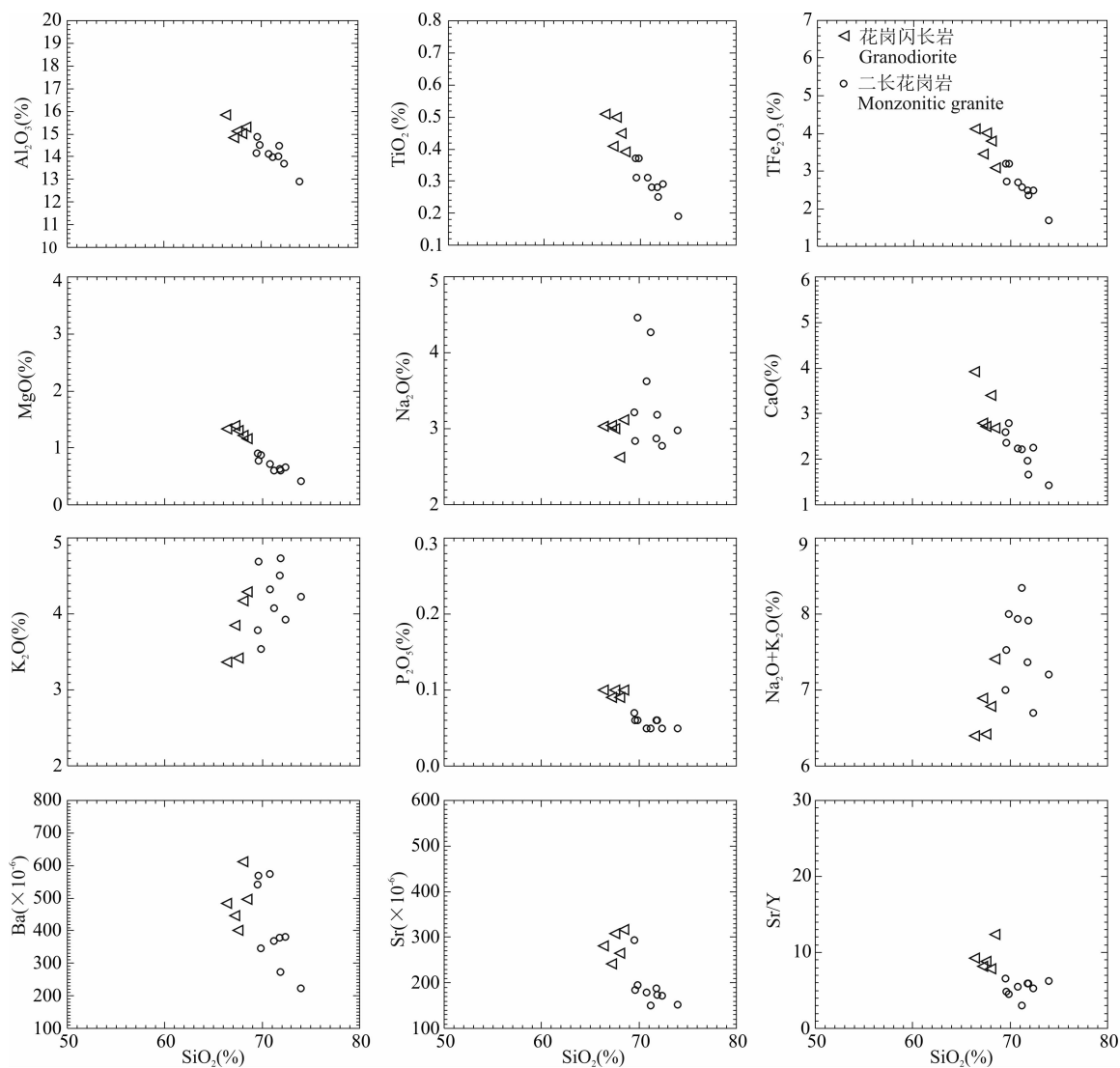


图 7 当家寺花岗岩体主要氧化物、微量元素 Harker 图解
Fig. 7 Harker plots of selected major and trace elements for the Dangjiashi granitic complex

岩、闪长岩、花岗闪长岩和二长花岗岩等岩石系列。如前述,对于研究区印支早期的构造背景及花岗岩的成因尚存在争议,但多数认为与海西—印支期南北板块的汇聚碰撞有关(Jin Weijun et al., 2005; Zhang Hongfei et al., 2006; Zhang Chengli et al., 2008; Yin Yong et al., 2009; Li Ting et al., 2012; Xu Xueyi et al., 2012; Guo et al., 2012; Luo Biji et al., 2012; Wei Ping et al., 2013; Jin Xiaoye et al., 2013; Huang Xiongfei et al., 2014)。本课题组获得区内江西沟花岗岩体南侧的拉木陇哇辉长岩侵位年龄为 247.8 ± 1.2 Ma, 岩石具有典型的弧岩浆特征(裴先治等, 未刊资料), 江西沟花岗岩体东南侧形成于早三叠世晚期—中三叠世早期的沟后杂岩体的辉长岩、辉长闪长岩、石英闪长岩和花岗闪长岩岩石组合和先后侵位关系(张永明等, 待刊资料), 与岛弧及大陆边缘弧火成

岩组合(Deng Jinfu et al., 2007, 2015a)类似;通过对区内隆务河组(T_1l)砂岩的地球化学分析,认为其源区构造背景为活动大陆边缘或大陆岛弧(裴先治等, 未刊资料),暗示青海南山构造带印支早期存在洋壳俯冲消减相关的弧环境。当家寺花岗岩体形成于中三叠世,岩石富集大离子亲石元素和亏损等高场强元素,可能反映了俯冲环境的影响。在 Rb-Yb+Ta 图解上(图 10a),当家寺花岗岩体样品点落入火山弧花岗岩(VAG)与同碰撞花岗岩(Syn-COLG)界线区域; Rb-(Y+Nb)图解上(图 10b)样品点落入火山弧花岗岩(VAG)、同碰撞花岗岩(Syn-COLG)和板内花岗岩(WPG)的界线附近区域。上述特征表明当家寺花岗岩体兼具岛弧花岗岩和同碰撞花岗岩的特征,显示构造转换体制下花岗岩的地球化学特征,可能形成于俯冲阶段的末期。

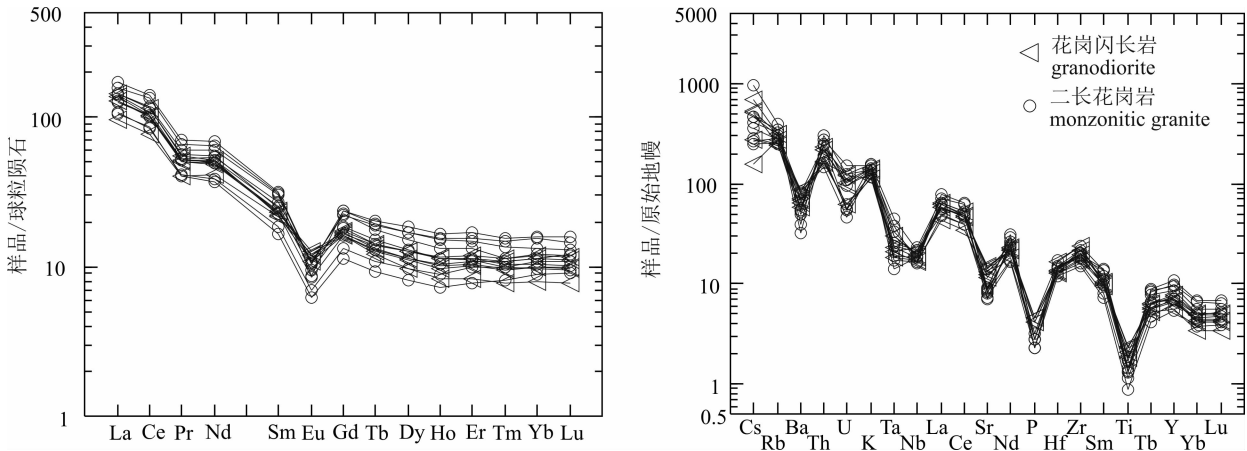


图 8 当家寺花岗岩体稀土元素球粒陨石标准化配分曲线图(a)(球粒陨石标准化数据值引自 Boynton,1984)和微量元素原始地幔标准化蛛网(b)(原始地幔标准化数据值引自 Sun and McDonough,1989)

Fig.8 Chondrite-normalized REE patterns (a) (normalization values after Boynton, 1984) and primitive-mantle normalized trace element spider diagrams (b) (normalization values after Sun and Mcdonough,1989) for Dangjiasi granitic complex

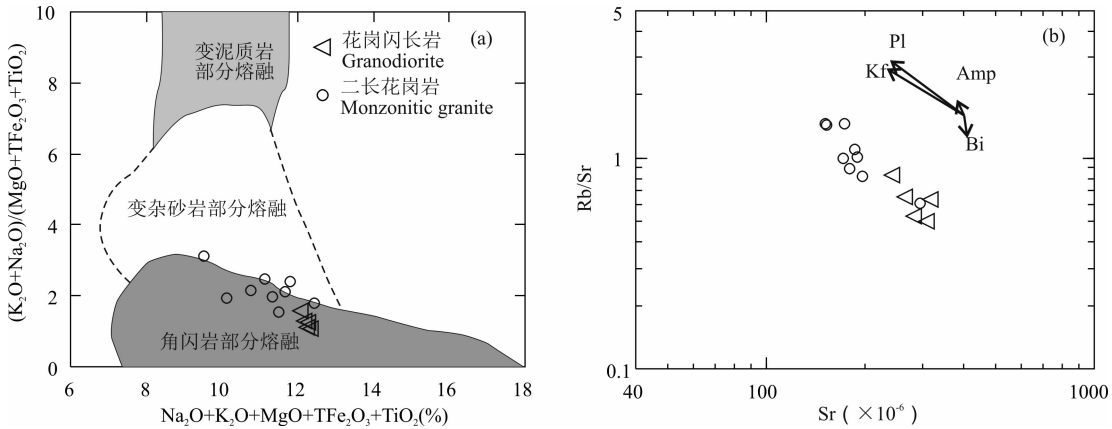


图 9 当家寺花岗岩体岩石化学成分与变泥质岩、变杂砂岩、变角闪岩派生的实验熔体化学成分对比 (a,原图据 Kaygusuz et al. ,2008)和 Rb/Sr-Sr 图解(b)

Fig.9 Chemical contents contrast between Dangjiasi granitic complex and meta-mudstone-,greywacke-, meta-amphibolite-derived experimental melt (a,after Kaygusuz et al. ,2008) and Rb/Sr-Sr diagram (b)

4.3 构造意义

区域上,东昆仑造山带、柴北缘构造带泥盆系牦牛山组磨拉石组合标志着早古生代造山运动的结束(Pan et al. , 1996; Li Rongshe et al. , 2007; Xu Zhiqin et al. ,2007; Lu Lu et al. ,2010),也拉开了古特提斯洋形成的序幕。石炭纪—二叠纪期间,从柴达木地块北缘(包括宗务隆构造带)至柴达木地块以东的西秦岭造山带西部都处于持续拉伸状态。宗务隆构造带晚古生代蛇绿岩的出现(Wang Yizhi et al. ,2001)标志着宗务隆构造带于晚石炭世拉开成有限小洋盆,该洋盆与阿尼玛卿晚古生代洋盆的发育时代基本一致(Bian et al. ,2004; Yang Jingsui et al. ,2004),后经历了洋盆的向南俯冲作用,形成了

天峻南山花岗岩体(246 Ma)、青海湖南山花岗岩体(238 Ma)、晒勒克郭来花岗岩体(249.2±2.6 Ma)和察汗诺花岗岩体(242.7±1.9 Ma 和 243.5±2.4 Ma)等俯冲型花岗岩(Guo Anlin et al. ,2009; Peng Yuan et al. ,2016),后碰撞期的局部拉伸形成了相关基性岩墙群和二郎洞 A 型花岗岩体(215 Ma)(Guo Anlin et al. ,2009)。青海南山构造带内发育的石炭—二叠系和下三叠统指示该构造带晚古生代—早中生代的沉积环境与宗务隆构造带存在较强的空间上的延续性。Guo Anlin et al. (2009)也认为宗务隆构造带的晚古生代和早—中三叠世沉积建造向东可能经橡皮山、青海南山与西秦岭商丹带晚古生代和早—中三叠世残余海盆相连。当家寺花岗

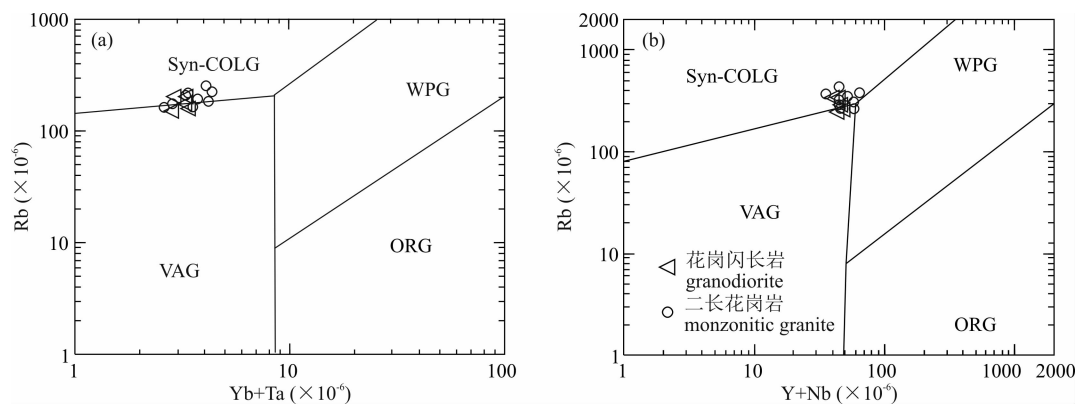


图 10 当家寺花岗岩体 Rb-(Yb+Ta)图解(a)和 Rb-(Y+Nb)图解(b)(据 Pearce,1996)
Fig. 10 Tectonic setting diagrams Rb-(Yb+Ta) (a) and Rb-(Y+Nb) (b)
for Dangjiashi granitic complex (after Pearce,1996)

岩体形成于中三叠世(241~240 Ma),其具有活动大陆边缘弧花岗岩的地球化学特征,表明研究区在中三叠世处于洋盆俯冲背景下的活动大陆边缘构造环境。Hao Taiping(1990)对龙羊峡地区中酸性火山岩研究认为其成岩环境也为活动大陆边缘,与同期侵入岩为同源岩浆演化。上述资料指示青海南山构造带可能为宗务隆构造带的东延,即存在晚古生代一早中生代有限小洋盆。在洋盆俯冲过程中形成了中三叠世青海南山构造带内广泛的基性—中酸性岩浆侵入,青海南山构造带印支期岩浆岩带分布位置显示其可能为南祁连地块与西秦岭地块间的碰撞并向南俯冲的结果,但有学者(Sun Yangui,2004)认为其为西秦岭地块沿共和盆地北缘(早中三叠世时期共和拗拉谷北缘谷坡与深水盆地的过渡地带)向北俯冲并与南祁连地块在深部发生碰撞的岩石学记录。区域上,在同仁-尖扎一带断续出露有晚二叠世—早三叠世超镁铁质-镁铁质岩石,可能指示西秦岭北缘存在一条晚二叠世—早三叠世蛇绿混杂岩带(Zhang Kexin et al.,2007;Wang Huiqing et al.,2009,2010)。同时,同仁-夏河一带发育有二叠纪具岛弧火山岩地球化学特征的火山-岩浆作用(Kou Xiaohu et al.,2007);龙羊峡地区也出露中—晚三叠世具弧岩浆岩特征的火山岩(Hao Taiping,1990)及密切相关的深成侵入岩(当家寺花岗岩体),该火山-岩浆弧可能与天峻南山共同形成一条位于祁连造山带南缘的弧火山-岩浆带。Yan Zhen et al.(2012)认为该弧火山-岩浆带与其北侧早古生代弧火山-岩浆带在时间上呈现出向南跃迁特征,并于三叠纪向北拼贴于南祁连南缘。但深地震反射剖面资料(Gao Rui et al.,2006;Wang Haiyan et al.,

2014)显示西秦岭北侧下地壳具南倾的反射特征,其揭示了扬子与华北两个大陆板块在西秦岭造山带北缘的汇聚事件。该汇聚事件可能包含了印支期宗务隆洋的向南俯冲过程的信息。因此,当家寺花岗岩体可能是在宗务隆有限洋盆向南俯冲背景下由俯冲流体交代的上覆地幔楔熔融形成基性岩浆底侵下地壳,造成基性下地壳发生角闪石脱水部分熔融形成的。同时由于在源区有幔源基性岩浆的注入,并发生一定程度的岩浆混合作用,使得长英质岩浆成分受到轻微改造,未完全混合的基性岩浆形成了暗色微粒包体。

5 结论

- (1)当家寺花岗岩体的花岗闪长岩和二长花岗岩 LA-ICP-MS 锆石 U-Pb 年龄分别为 240.1 ± 2.1 Ma 和 241.0 ± 2.6 Ma,代表岩体形成于中三叠世。
- (2)岩石具有高硅、富钾和准铝质—弱过铝质特征,属高钾钙碱性 I 型花岗岩,可能由存在于青海南山构造带内的晚古生代—早中生代有限小洋盆的俯冲造成上覆地幔楔熔融形成基性岩浆底侵下地壳,使得基性下地壳部分熔融而形成,其源区受到一定程度的幔源岩浆的混染。
- (3)青海南山构造带可能为宗务隆构造带的东延,即研究区存在晚古生代—早中生代有限小洋盆,当家寺花岗岩体是其在在中三叠世向南俯冲末期的产物。

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LA-ICP-MS Zircon U-Pb Dating and Geochemistry of the Dangjiasi Granitic Complex in the Qinghai Nanshan Tectonic Zone, and Its Geological Implications

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Abstract

An NW-SE trending Indosinian granitoid belt occurs within the Qinghai Nanshan tectonic zone, and the study of this granite zone has great scientific significance for investigating the transitional relation between the West Qinling and the Qilian orogenic belts in Paleozoic Tethys evolution stage. The Dangjiasi granitic complex consists dominantly of granodiorite pluton and monzogranite pluton. In this paper, a detailed study for the LA-ICP-MS zircon U-Pb dating, petrology and geochemistry characteristics was carried out for the Dangjiasi granitic complex. The results show that the crystallization ages of the Dangjiasi granodiorite and monzogranite are 240.1 ± 2.1 Ma and 241.0 ± 2.6 Ma, respectively, belonging to the Middle Triassic. The whole rock geochemical data show that the Dangjiasi granitic complex is relatively rich in SiO_2 (66.37%~73.99%) and K_2O (3.37%~4.73%), suggesting that the granitic rocks are metaluminous to weak peraluminous high-K calc-alkaline I-type granite. The rocks are enriched in LREE and depleted in HREE. The REE patterns show rightward incline and moderate negative Eu anomaly ($\delta\text{Eu}=0.32 \sim 0.64$). The trace element geochemistry is evidently characterized by positive anomaly of LILE (e. g. Cs, Rb, K) and negative anomaly of HFSE (e. g. Nb, Ta, P, Ti). The Dangjiasi granitic complex originated mainly from the partial melting of lower crustal basic rocks, with minor mantle-derived component. In combination with analyses of regional geological setting, we suggest that the Dangjiasi granitic complex was formed in the stage of the southward subduction of the Zongwulong Oceanic crust.

Key words: Qinghai Nanshan; northern margin of West Qinling; Dangjiasi granitic complex; LA-ICP-MS zircon U-Pb dating; geochemistry; tectonic setting