

内蒙古赤峰南窝铺铀矿床英安岩的 SHRIMP 锆石 U-Pb 年龄、地球化学特征及地质意义



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内容提要: 南窝铺铀矿床位于西拉木伦缝合带以南、康宝—围场—赤峰断裂带(华北北缘断裂带的一部分)以北的白乃庙加里东期岛弧带和华北古板块北缘海西期俯冲—碰撞带, 属沽源—红山子铀成矿带北东段红山子—广兴铀成矿亚带的一个铀矿床, 铀矿体赋存在额里图组安山岩—英安岩—流纹岩组合中。该组合中的英安岩 SHRIMP 锆石 $n(^{206}\text{Pb})/n(^{238}\text{U})$ 加权平均年龄为 268.1 ± 2.5 Ma ($N = 15$, MSWD = 0.85), 结合安山质角砾凝灰岩的锆石 U-Pb 年龄为 277.1 ± 0.9 Ma, 指示额里图组安山岩—英安岩—流纹岩的地质时代属早二叠世(乌拉尔世)—中二叠世(瓜德鲁普世)早期; 英安岩 $\text{SiO}_2 = 65.49\% \sim 68.23\%$, $\text{K}_2\text{O} = 1.83\% \sim 2.56\%$, $\text{Na}_2\text{O} = 4.13\% \sim 5.01\%$, $(\text{K}_2\text{O} + \text{Na}_2\text{O}) = 5.96\% \sim 7.47\%$, $\text{K}_2\text{O}/\text{Na}_2\text{O} = 0.43 \sim 0.59$, $(\text{FeO} + \text{Fe}_2\text{O}_3) = 3.13\% \sim 4.63\%$, $\text{MgO} = 1.25\% \sim 1.79\%$, 在 TAS 图解上落入亚碱性系列英安岩区, 在 FAM 图解上落入钙碱性系列范围, 在 $\text{SiO}_2 - \text{K}_2\text{O}$ 图解上落入中钾钙碱性系列范围, 在 $\text{Na}_2\text{O} - \text{K}_2\text{O}$ 图解上落入 I 型花岗岩区; $\text{Al}_2\text{O}_3 = 15.74\% \sim 16.77\%$, $\text{CaO} = 2.87\% \sim 3.59\%$, $\text{A/CNK} = 0.95 \sim 1.12$ (平均为 1.04, <1.1), 标准矿物刚玉(C)的含量为 0~1.74%, 平均为 0.96%, 指示源岩为 I 型岩浆岩; 英安岩稀土元素总量低, 富集轻稀土, Eu 无明显负异常, $\Sigma\text{REE} = 76.5 \times 10^{-6} \sim 95.4 \times 10^{-6}$, $(\text{La/Yb})_N = 4.97 \sim 12.5$ (平均 8.95), $\delta\text{Eu} = 0.80 \sim 1.13$ (平均 0.94), 稀土配分模式为右倾型, 与安第斯型钙碱性系列火山岩基本一致。英安岩明显富集大离子亲石元素 Rb、Th、U、K、Sr 等和亏损高场强元素 Nb、Ta、P、Ti 等, 微量元素蛛网图与安第斯型钙碱性系列火山岩的形式一致, 在 $\text{Ce} - \text{SiO}_2$ 、 $\text{Al}_2\text{O}_3 - \text{Ga}$ 图解上均落入 I 型花岗岩区, 在 $\text{Ta} - \text{Yb}$ 、 $\text{Nb} - \text{Y}$ 构造环境判别图解上落入火山弧花岗岩区域(VAG)及同碰撞花岗岩(syn-COLG)交界处, 在 $\text{Rb} - (\text{Y} + \text{Nb})$ 、 $\text{Rb} - (\text{Yb} + \text{Ta})$ 构造环境判别图解上落入火山弧花岗岩区域(VAG), 指示英安岩形成于俯冲作用下的岛弧环境。南窝铺铀矿床英安岩地质时代、岩石系列和形成构造环境的确定, 不仅证实了早二叠世—中二叠世早期额里图组是红山子—广兴铀成矿亚带一个新的赋矿层位, 而且揭示了岛弧型安山岩—英安岩—流纹岩组合也赋存有与火山岩有关的热液型铀矿, 为深入开展热液型铀矿成矿理论的研究和进一步扩大铀矿勘查范围提供了新的基础资料。

关键词: 英安岩; 下二叠统—中二叠统; 额里图组; 内蒙古赤峰南窝铺铀矿床

南窝铺铀矿床位于沽源—红山子铀成矿带北东段的红山子—广兴铀成矿亚带, 铀矿体赋存在额里图组安山岩—英安岩—流纹岩组合中。大致以康宝—围场—赤峰断裂带(华北北缘断裂带的一部分)为界, 沽源—红山子铀成矿带可分为北东段的红山子—广兴铀成矿亚带和西南段的沽源—丰宁铀成矿亚带, 与火山岩有关的热液型铀矿床的赋矿火山岩在红山子—广兴铀成矿亚带曾被归于晚侏罗世满克头鄂博组、玛尼吐组和白音高老组, 而在沽源—

丰宁铀成矿亚带被归于早白垩世张家口组, 虽然两个成矿亚带的赋矿火山岩被归于不同的岩石地层单位, 但它们的岩石组合和地质时代长期被认为是晚侏罗世流纹岩—粗面岩组合(河北省地质矿产局, 1996; 内蒙古自治区地质矿产局, 1996)。近十年来, 随着基础地质研究的不断深入和铀矿勘查工作的持续开展, 现已查明这两个成矿亚带的赋矿火山岩不仅岩石组合不同, 而且地质时代也不一致。红山子—广兴铀成矿亚带与火山岩有关的热液型铀矿床

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的赋矿火山岩主要属于晚侏罗世新民组高钾钙碱性流纹岩—碱性流纹岩组合,流纹岩具有 A 型流纹岩的特征,是板内拉张构造环境下正常厚度地壳的年轻下地壳部分熔融的产物(巫建华等,2013,2016,2017a,2017b; 彭啟辉,2015; 解开瑞等,2016; 黎伟等,2017; 姜山等,2018);而沽源—丰宁铀成矿亚带的赋矿火山岩主要属于早白垩世张家口组高钾钙碱性流纹岩—碱性粗面岩组合,流纹岩和粗面岩分别具有 A 型流纹岩和高压型粗面岩的地球化学特征,分别是板内拉张构造环境下地壳中上部和底部部分熔融的产物(邓晋福等,2000; 巫建华等,2014,2015,2017b,2017c; 夏应冰等,2016; 张雅菲等,2016; 林天发等,2019)。同时,在红山子—广兴铀成矿亚带参照红山子铀矿床、在沽源—丰宁铀成矿亚带参照张麻井铀矿床开展铀矿勘查工作,取得了较好的找矿效果,并在红山子—广兴铀成矿亚带不整合于新民组高钾钙碱性流纹岩—碱性流纹岩组合之下的额里图组安山岩—英安岩—流纹岩组合中发现了南窝铺铀矿床和众多铀矿点、铀异常点(祝洪涛等,2014; 纪宏伟,2015; 黎伟等,2017; 巫建华等,2017b)。然而,对赋存南窝铺铀矿床和众多铀矿点、铀异常点的额里图组安山岩—英安岩—流纹岩组合却缺乏系统的年代学和地球化学研究,制约了红山子—广兴铀成矿亚带铀矿勘查的开展和铀成矿理论的完善。本文以南窝铺铀矿床赋矿安山岩—英安岩—流纹岩组合中的英安岩为研究对象,通过 SHRIMP 锆石 U-Pb 定年和主、微量元素研究,查明其地质时代、岩石系列和形成的构造环境,并对其地质意义进行探讨。

1 区域地质背景

红山子—广兴铀成矿亚带位于西拉木伦河—长春缝合带以南、康宝—围场—赤峰—开原断裂带以北的华北古板块北缘,加里东构造阶段属白乃庙岛弧带—温都尔庙—翁牛特旗增生杂岩带,海西构造阶段属华北古板块北缘俯冲岛弧带,印支构造阶段早期属华北古板块北缘碰撞带,印支构造阶段晚期晚三叠世华北古板块北缘伸展带(图 1a)。寒武纪—志留纪末期,华北克拉通北部被动大陆边缘与白乃庙岛弧带之间为宽广的洋盆(赵越等,2010; 徐备等,2014),白乃庙岛弧带发育火山沉积岩系(Jian Ping et al., 2008; 张维等,2008; 李锦轶,2009),主要由绿片岩相—低角闪岩相变质沉积岩、火山岩和侵入岩组成(Zhang Shuanhong et al., 2014);晚志留世—早泥盆世,古亚洲洋向南俯冲,白乃庙岛弧北侧

发育温都尔庙—翁牛特旗增生杂岩带、南侧以弧—陆软碰撞的方式拼贴在华北克拉通北缘,并形成了西别河组磨拉石或类磨拉石组合(许立权等,2003; 王平,2005; Chen Xiuqin and Boucot, 2007; 张允平等,2010)。泥盆纪期间,在华北北缘发育有与弧—陆碰撞后伸展有关的碱性杂岩及双峰式火山岩(Zhang Qiqi et al., 2018)。海西构造阶段,古亚洲洋继续向南俯冲,华北古板块北缘发育一系列与安第斯型大陆边缘弧相关的石炭纪—二叠纪岩浆岩(江小均等,2011; Zhang Shuanhong et al., 2016)和火山岩(曹花花等,2012; 曹代勇等,2014; 董晓杰等,2016; 彭斌等,2016; 王月古,2018; 崔玉良等,2019)。集宁北部商都县西井子镇一带早—中二叠世(267~272 Ma)苏吉组酸性火山岩和赤峰北部朝阳沟村早二叠世流纹岩具有安第斯型活动大陆边缘火山岩的地球化学特征(董晓杰等,2016; 崔玉良等,2019),赤峰铭山隆起带早二叠世额里图组安山岩形成的构造环境为火山弧(彭斌等,2016),赤峰翁牛特旗早二叠世流纹岩、安山岩和流纹质—英安质凝灰岩形成于与俯冲带有关的岛弧或活动大陆边缘的构造环境(曹代勇等,2014),吉林中部早二叠世大河深组流纹岩—英安岩—粗面英安岩组合显示活动大陆边缘的构造背景,延吉地区早二叠世关门咀子组玄武安山岩形成于岛弧构造环境(曹花花等,2012)。印支—燕山构造阶段,古亚洲洋关闭,华北古板块与西伯利亚古板块最后沿阴山—燕山造山带一线拼贴,形成狼山蛇绿混杂岩(吕洪波等,2018)。中生代晚期,红山子—广兴铀成矿亚带受太平洋板块向北西俯冲的影响,该区域进入后碰撞造山系统,由挤压构造体制进入伸展造山构造体制,侏罗纪、白垩纪岩浆活动强烈(孟恩等,2011; 许文良等,2013; Tang Jie et al., 2015),形成了晚侏罗世早期新民组火山岩岩系、侵入岩和早白垩世侵入岩(巫建华等,2013,2016,2017a; 解开瑞等,2016; 姜山等,2018; 王常东等,2019; 祝洪涛等,2019,2020)。古亚洲洋构造成矿域与环太平洋构造成矿域的叠加、复合和转换,使研究区成矿地质条件优越,有色金属、稀有金属矿产集中分布(苏美霞等,2020)。

2 矿床地质特征

南窝铺铀矿床位于红山子—托河复式岩体隆起带南东侧,矿床受大兴永—南窝铺断裂控制,矿区内出露的地层主要为下二叠统—中二叠统额里图组、中二叠统于家北沟组、上侏罗统新民组和中新统汉

诺坝组(图1b)。额里图组为一套陆相火山岩、沉积碎屑岩,下部为长石砂岩、砂岩、粉砂质页岩、粉砂岩,上部为灰—灰褐色英安岩、紫色、灰—灰褐色安山岩、安山质凝灰角砾岩、紫红—褐色熔结凝灰岩、流纹岩、沉凝灰岩夹长石砂岩;于家北沟组为一套海陆交互相岩石组合,岩性主要为黄绿色凝灰质砂砾岩、砂岩、凝灰岩、安山岩。新民组为一套高钾钙碱性流纹岩—碱性流纹岩组合,主要有流纹岩、流纹斑岩、流纹质熔结凝灰岩、流纹质凝灰岩和火山角砾岩

组成,流纹岩 SHRIMP 钨石 U-Pb 年龄为 156~158 Ma(巫建华等,2013,2017a;解开瑞等,2016;姜山等,2018)。中新统汉诺坝组为玄武岩。侵入岩主要由广兴海西期花岗闪长岩和红山子—托河燕山期复式岩体组成。广兴花岗闪长岩 SHRIMP 钨石 U-Pb 年龄为 263.3±2.5 Ma(江小均等,2011);红山子复式岩体粗粒碱长花岗岩和斑状黑云母花岗岩的 LA-ICP-MS 钨石 U-Pb 年龄分别为 154±1 Ma 和 151±1 Ma,细粒黑云母碱长花岗岩和花岗斑岩的 LA-

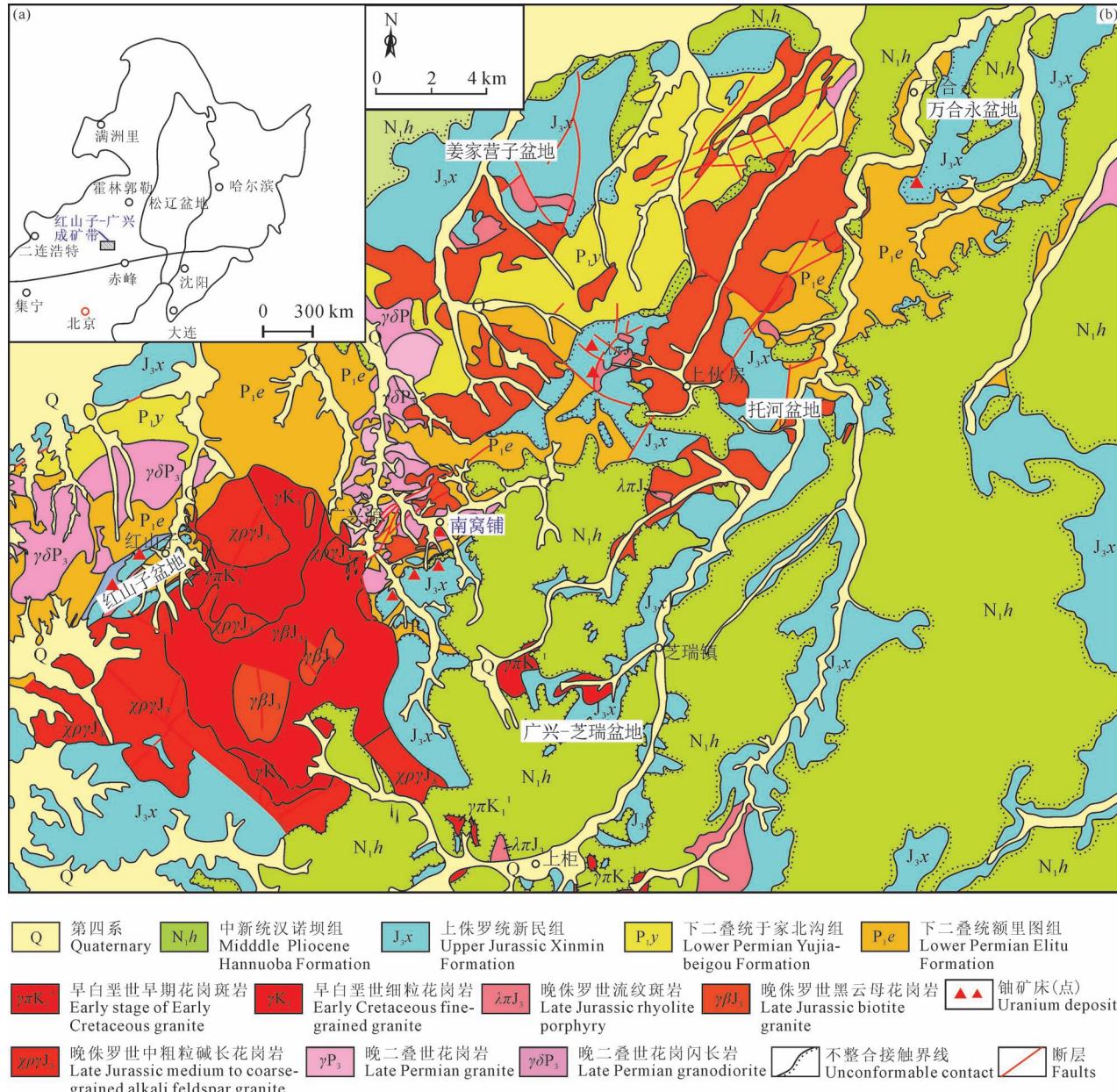


图 1 内蒙古赤峰红山子—广兴铀成矿亚带大地构造位置(a,据丁辉等,2016 修改)
和地质简图(b,据祝洪涛等,2019 修改)

Fig. 1 Tectonic location map (a, modified from Ding Hui et al., 2016&) and geological map (b, modified from Zhu Hongtao et al., 2019&) of the Hongshanzi—Guangxing uranium metallogenetic subzone, Chifeng area, Inner Mongolia

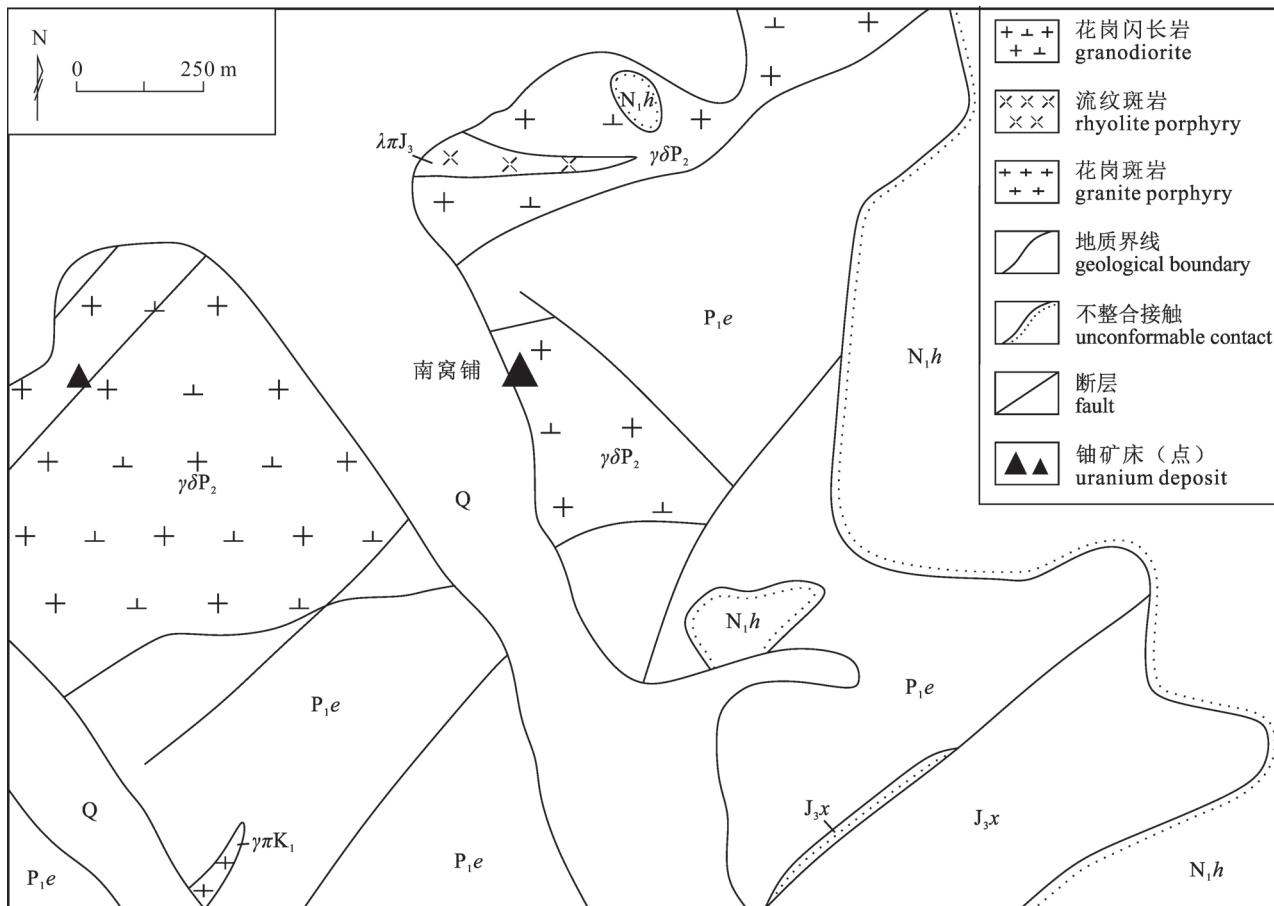


图 2 内蒙古赤峰南窝铺铀矿床地质简图

Fig. 2 Simplified geological map of the Nanwopu uranium deposit, Chifeng area, Inner Mongolia

Q—第四系;N₁h—中新统汉诺坝组;J₃x—上侏罗统新民组;P₁e—下二叠统额里图组;γπK₁—早白垩世花岗斑岩;λπJ₃—晚侏罗世流纹斑岩;γδP₂—中二叠世花岗闪长岩 Q—Quaternary; N₁h—Pliocene Hannuoba Formation; J₃x—Upper Jurassic Xinmin Formation; P₁e—Lower Permian Elitu Formation; γπK₁—Early Cretaceous granite porphyry; λπJ₃—Late Jurassic rhyolite porphyry; γδP₂—Middle Permian granodiorite

ICP-MS 锆石 U-Pb 年龄分别为 132 ± 2 Ma 和 133 ± 1 Ma(祝洪涛等, 2019);侵入红山子盆地新民组的花岗斑岩 SHRIMP 锆石 U-Pb 年龄为 134.8 ± 0.9 Ma(丁辉等, 2016)。

南窝铺铀矿床矿体产于断裂带 0~100 m 范围内,赋存在额里图组和海西期花岗闪长岩中(图 2),矿体长 100~230 m、沿倾向方向延伸 50~160 m,形态多呈透镜状,部分为似层状,产状 $150^\circ \sim 170^\circ \angle 70^\circ$,平面上侧列,剖面上斜列,向南西侧伏,侧伏角 20°左右。在 25 号矿点新发现的 4 个工业铀矿孔(ZK25-1、ZK25-3、ZK25-5、ZKNW5-1)和 1 个矿化孔(ZKNW4-1)赋存在额里图组英安质晶屑凝灰岩、英安岩(图 3a)中;ZK11-27 钻孔在孔深 110~410 m 范围内揭露到工业铀矿化、铀矿化和铀异常段赋存在花岗闪长岩中(图 3b)。

3 岩相学特征

本次研究的英安岩样品取自南窝铺铀矿床的 ZK25-3 钻孔 113.5~116.5 m 处,岩石呈灰白色、浅灰色,块状构造(图 4a),斑状结构,斑晶由斜长石、少量钾长石和石英、暗色矿物假像构成,大小一般 0.25~2.0 mm,杂乱分布。其中斜长石呈近半自形板状,隐约可见聚片双晶,部分见绢云母化、碳酸盐化和黝帘石化,表面很“脏”,局部可见聚斑;钾长石呈近半自形板状,表面裂纹发育,部分被高岭石、碳酸盐交代;石英呈他形粒状,含量较少;暗色矿物呈他形粒状,被铁质、绿帘石交代呈假像(图 4b)。基质主要由长英质、不透明矿物和少量玻璃质构成包含霏细结构,大小一般 0.03~0.1 mm,界线清楚,基质中不透明矿物和玻璃质星散状分布。

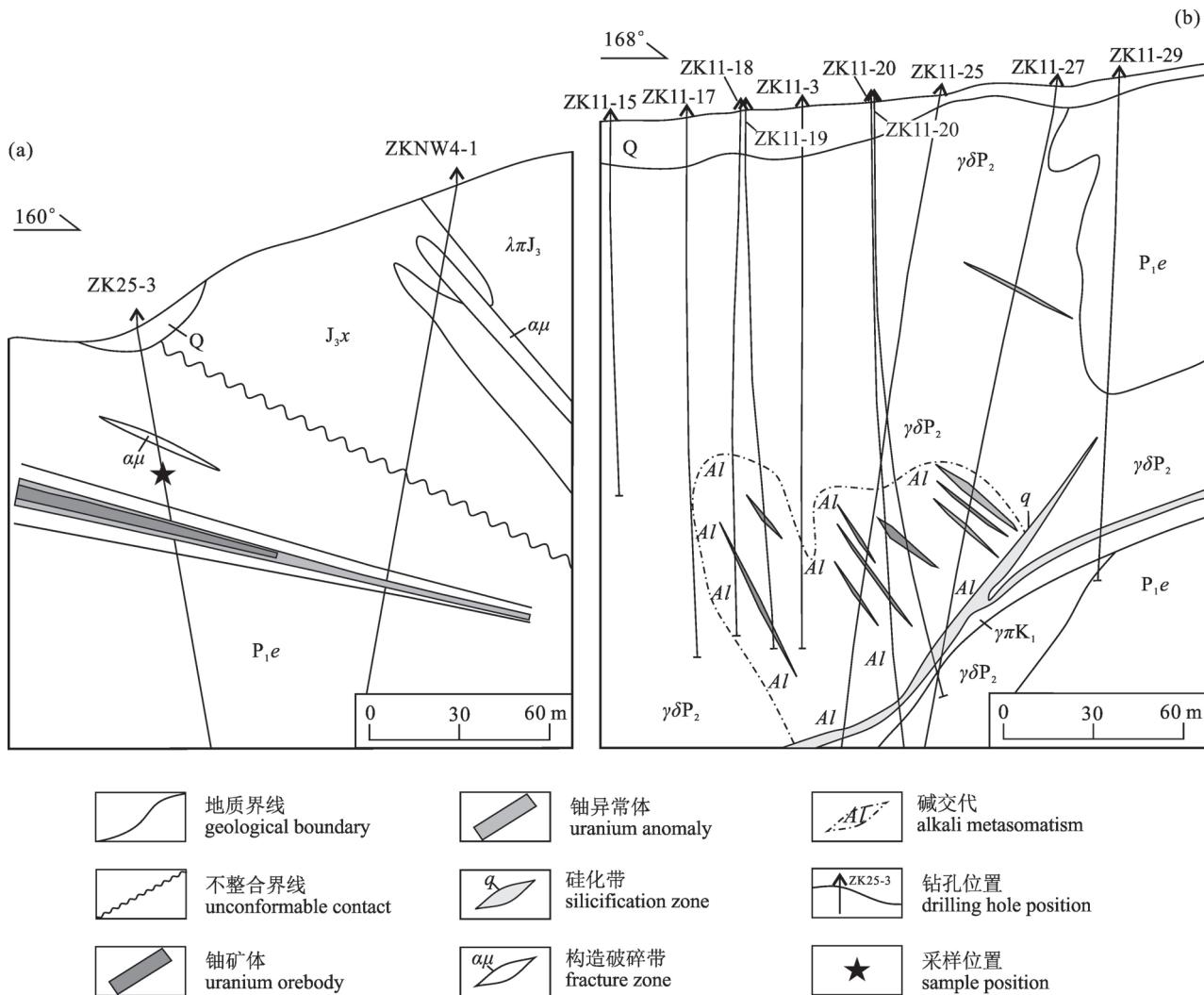


图3 内蒙古赤峰南窝铺铀矿床 NW4(a)、11号线(b)剖面图

Fig. 3 The Profiles of NW4 (a) and the Line 11 (b) of the Nanwopu uranium deposit, Chifeng area, Inner Mongolia

Q—第四系; J_3x —上侏罗统新民组; P_1e —下二叠统额里图组; $\gamma\pi K_1$ —早白垩世花岗斑岩; $\lambda\pi J_3$ —晚侏罗世流纹斑岩; $\gamma\delta P_2$ —中二叠世花岗闪长岩 Q—Quaternary; J_3x —Upper Jurassic Xinmin Fromation; P_1e —Lower Permian Elitu Fromation; $\gamma\pi K_1$ —Early Cretaceous granite porphyry; $\lambda\pi J_3$ —Late Jurassic rhyolite porphyry; $\gamma\delta P_2$ —Middle Permian granodiorite

英安岩遭受到蚀变影响大多数发生绿泥石、绿帘石化,但2组节理夹角仍明显保留,边缘暗化边发育明显,主要为石英、磁铁矿等矿物环绕;黑云母斑晶呈半自形板状,多色性明显,溶蚀较强发育暗化边,粒径长度集中于0.3~0.5 mm。基质约占83%,多为隐晶质或蚀变成石英、长石微晶、磁铁矿等,玻璃质大多发生脱玻化成隐晶质或石英、长石微晶。

4 分析方法

4.1 SHRIMP 锌石 U-Pb 分析

锌石挑选在河北省廊坊市诚信地质服务有限公司完成,火山岩样品取10 kg,破碎至80~120目,洗

去粉尘,经淘洗除去轻矿物,保留重矿物,再用永久磁铁除去磁铁矿等强磁性矿物,经重液分选除去比重小于锌石的矿物,最后在双目镜下人工精选出锌石晶体。将挑选好的锌石与标准锌石 Temora(年龄为417 Ma)一起粘贴,制成环氧树脂样品靶。干燥后,打磨、抛光使锌石中心部分暴露,然后进行反射光、透射光和阴极发光显微照相。反射光、透射光和阴极发光显微照相在中国地质科学院矿产资源研究所完成,锌石 SHRIMP U-Pb 分析在北京离子探针中心 SHRIMP-II 上完成。年龄测试前,利用反射光、透射光显微照片选择无裂纹、无包裹体、表面洁净的晶体,利用阴极发光照片选择具有明显环带结构的

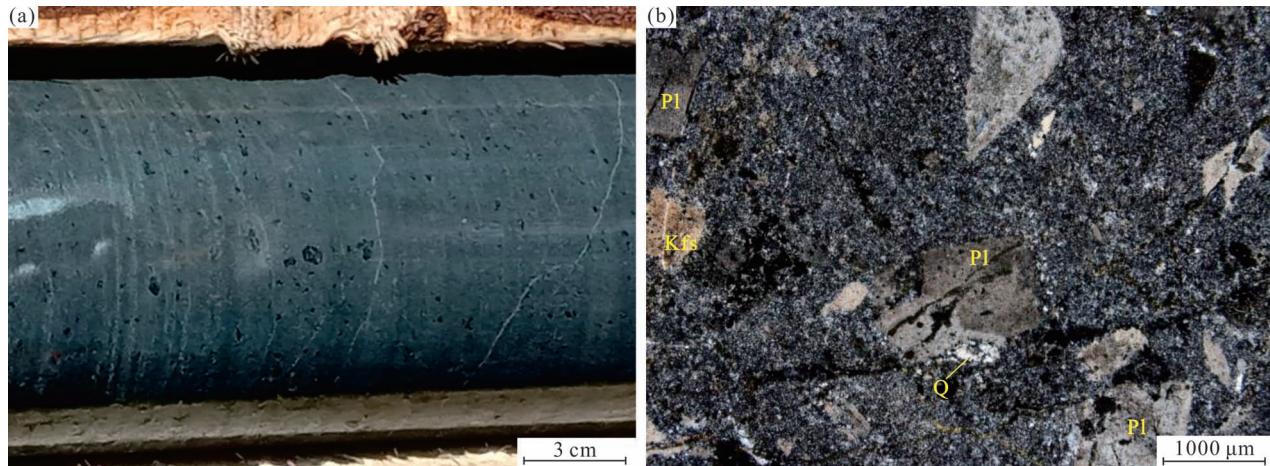


图 4 内蒙古赤峰地区南窝铺铀矿床英安岩(a)钻孔照片和(b)镜下照片

Fig. 4 (a) Borehole photographs and (b) microscopic photographs of dacite in Nanwopu uranium deposit, Chifeng area, Inner Mongolia
Pl—斜长石;Kfs—钾长石;Q—石英
Pl—plagioclase;Kfs—K-feldspar;Q—quartz

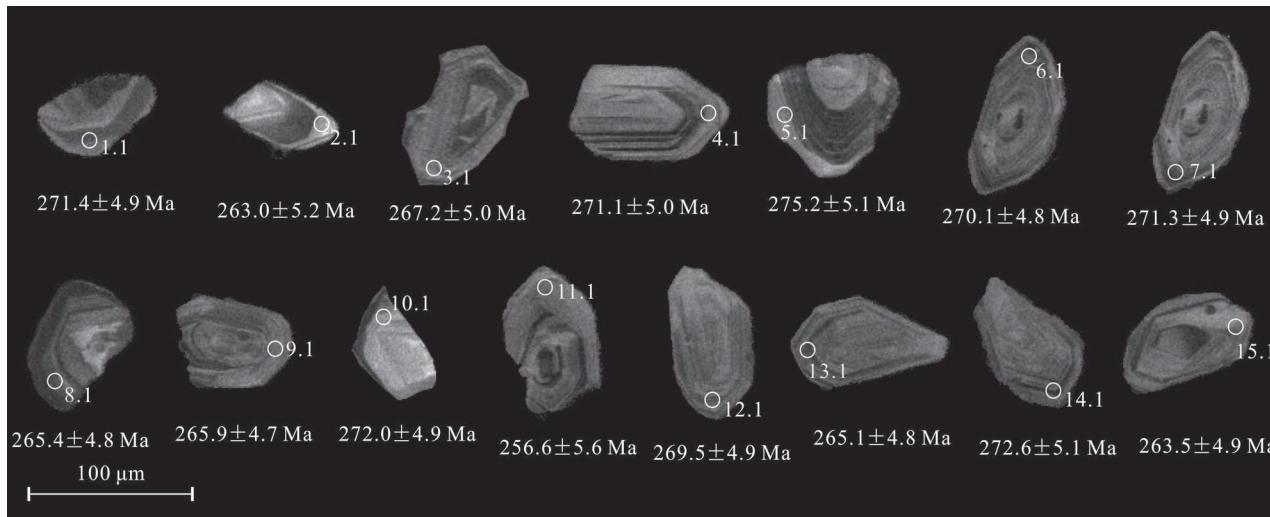


图 5 内蒙古赤峰南窝铺铀矿床英安岩锆石阴极发光图像、SHRIMP 分析点位及 $n(^{206}\text{Pb})/n(^{238}\text{U})$ 年龄值

Fig. 5 Cathodoluminescence photos, SHRIMP analytical spots and corresponding $n(^{206}\text{Pb})/n(^{238}\text{U})$ ages of dacite from the Nanwopu uranium deposit, Chifeng area, Inner Mongolia

岩浆成因锆石。待测锆石挑选之后,随机对待测锆石进行年龄测试。详细的分析流程和原理参见宋彪等(2002),数据处理、年龄计算采用 Ludwig 博士编写的 SQUID1.0 及 ISOPLOT 程序(Ludwig, 2003)。南窝铺铀矿床英安岩样品的 SHRIMP 锆石 U-Pb 分析结果列于表 1。

4.2 主量元素和微量元素分析

本文岩石薄片的制备和碎样委托河北省区域地质矿产调查研究所实验室完成,通过显微镜下鉴定之后确定全岩地球化学分析样品。主、微量元素分

析测试在核工业北京地质研究院分析测试中心完成。主量元素分析测试采用化学分析法(CA)和 X 射线荧光光谱法(XRF)。化学分析法主要分析氧化亚铁的含量,X 射线荧光光谱法在 AxiosMAX X 射线荧光光谱仪上完成,测试前的样片制作可参见周万蓬(2015),实验过程中,X 射线管电压为 50 kV,电流为 50 mA,元素分析相对误差小于 5%,检测方法和依据参照 GB/T 14506.14-2010《硅酸盐岩石化学分析方法第 14 部分:氧化亚铁量测定》,GB/T 14506.28-2010《硅酸盐岩石化学分析方法第 28

表 1 内蒙古赤峰南窝铺铀矿床英安岩 SHRIMP 锆石 U-Pb 同位素分析结果

Table 1 Analytical results of SHRIMP zircons U-Pb isotopes of dacite from the Nanwopu uranium deposit, Chifeng area, Inner Mongolia

测点号	$^{206}\text{Pb}_{\text{e}}$ (%)	元素含量($\times 10^{-6}$)			同位素比值			同位素年龄(Ma)			
		Pb^*	Th	U	$n(^{207}\text{Pb}^*)/n(^{206}\text{Pb}^*)$			$n(^{206}\text{Pb}^*)/n(^{235}\text{U})$			
					测值	$\pm\%$	测值	$\pm\%$	测值	$\pm\%$	
英安岩、南窝铺, NWP301											
1.1	-	7.00	143	190	0.78	0.05330	3.5	0.3160	3.9	0.04300	1.8
2.1	1.33	5.62	102	155	0.68	0.04609	8.8	0.2646	9.1	0.04164	2.0
3.1	1.22	6.05	120	164	0.75	0.04438	8.1	0.2590	8.4	0.04233	1.9
4.1	0.74	6.96	115	187	0.63	0.04696	6.7	0.2781	6.9	0.04295	1.9
5.1	0.21	8.47	175	226	0.80	0.05292	3.0	0.3182	3.5	0.04361	1.9
6.1	0.33	9.27	224	251	0.92	0.04977	7.1	0.2936	7.3	0.04279	1.8
7.1	0.32	7.99	184	216	0.88	0.05250	6.0	0.3111	6.3	0.04298	1.8
8.1	0.40	5.80	84	160	0.55	0.05450	4.1	0.3158	4.5	0.04203	1.9
9.1	0.52	7.59	144	209	0.71	0.04876	4.2	0.2831	4.6	0.04210	1.8
10.1	-	6.23	106	169	0.65	0.05578	2.9	0.3315	3.5	0.04310	1.8
11.1	2.14	2.87	37	81	0.47	0.03874	19.9	0.2169	20	0.04061	2.2
12.1	0.70	7.49	151	203	0.77	0.04737	7.2	0.2788	7.4	0.04269	1.8
13.1	0.30	5.71	91	158	0.59	0.05371	3.6	0.3109	4.1	0.04198	1.9
14.1	0.49	5.03	86	135	0.66	0.04796	5.3	0.2856	5.6	0.04319	1.9
15.1	0.58	5.33	96	148	0.67	0.04994	4.4	0.2873	4.8	0.04172	1.9

注: $^{206}\text{Pb}_{\text{e}}$ 和 $^{206}\text{Pb}^*$ 分别表示普通铅和放射性成因铅; 普通铅根据实测 ^{204}Pb 进行校正; 误差为 1σ 。

部分: 16 个主次成分量测定》, 岩石矿物分析《第四版 16.20 灼烧减量的测定》。微量元素分析测试采用电感耦合等离子质谱法 (ICP-MS), 样品溶液的配置过程可参见周万蓬 (2015), 分析测试是在 NexION 300D 等离子体质谱仪上完成, 工作温度控制在 20℃, 相对湿度保持在 27%, 当微量元素含量小于 10 $\mu\text{g/g}$ 时, 测试相对误差小于 10%, 当微量元素含量大于 10 $\mu\text{g/g}$ 时, 测试相对误差小于 5%, 测试方法和依据参照 GB/T 14506.30-2010《硅酸盐岩石化学分析方法第 30 部分: 44 个元素量测定》。南窝铺铀矿床英安岩样品的主量元素和微量元素分析结果及有关参数列于表 2。

5 分析结果

5.1 锆石 U-Pb 年龄

样品 NWP301 的锆石颗粒长度约 100~150 μm , 长宽比为 1.5~2.0, 自形短柱状或双锥状, 阴极发光图像显示锆石多具有清晰的韵律环带结构(图 5), 为典型岩浆结晶锆石的内部结构。NWP301 样品的 U 含量较高, 介于 81×10^{-6} ~ 251×10^{-6} 之间, Th/U 比值在 0.47~0.92 之间, 大于 0.4, 具有典型的岩浆锆石成分特征(Belousova et al., 2002; Rubatto, 2002), 15 个分析点的 $n(^{206}\text{Pb})/n(^{238}\text{U})$ 年龄数据在 256.6~275.2 Ma 之间, 在 U-Pb 谐和图(图 6)上集中分布, 加权平均年龄为 268.1±2.5 Ma, MSWD=0.85, 代表了火山岩的形成年龄。

5.2 地球化学特征

5.2.1 蚀变影响

南窝铺铀矿床处火山岩系发育, 热液蚀变强烈, 在挑选了新鲜的岩石样品后, 为保证能够有效地利用样品元素特征进行岩石分类、成因探讨(Rollinson, 1993), 首先以样品烧失量(LOI)为横坐标做 Harker 图解(图 7)评估岩浆演化和岩石形成过程中热液蚀变对元素

表 2 内蒙古赤峰南窝铺铀矿床英安岩主量元素(%)、微量元素($\times 10^{-6}$)和稀土元素($\times 10^{-6}$)分析结果及有关参数Table 2 Major elements (%), trace elements ($\times 10^{-6}$) and rare earth element ($\times 10^{-6}$) composition of dacite from the Nanwopu uranium deposit, Chifeng area, Inner Mongolia

点号	NWP 101	NWP 102	NWP 103	NWP 301	NWP 303	NWP 304	NWP 305	点号	NWP 101	NWP 102	NWP 103	NWP 301	NWP 303	NWP 304	NWP 305
SiO ₂	65.17	65.24	64.16	65.74	65.61	65.10	65.82	Er	1.21	0.91	1.12	1.78	1.34	1.24	1.03
TiO ₂	0.65	0.58	0.61	0.52	0.49	0.67	0.65	Tm	0.22	0.14	0.17	0.12	0.25	0.20	0.15
Al ₂ O ₃	15.96	16.13	16.77	15.89	16.37	16.05	15.74	Yb	1.16	0.95	1.24	2.00	1.67	1.42	1.16
Fe ₂ O ₃	3.25	3.70	2.49	2.39	2.75	2.39	1.42	Lu	0.18	0.12	0.16	0.19	0.15	0.15	0.10
FeO	1.59	1.35	1.39	1.24	1.66	1.94	1.79	Σ REE	90.8	85.5	95.4	76.9	88.0	76.5	80.9
MnO	0.04	0.04	0.05	0.05	0.04	0.05	0.05	Σ LREE	81.2	78.2	86.8	65.5	79.4	64.5	70.5
MgO	1.60	1.49	1.79	1.25	1.53	1.46	1.59	Σ HREE	9.59	7.32	8.63	11.4	8.53	12.0	10.4
CaO	2.97	2.87	3.17	3.59	3.08	3.15	3.19	Σ L/ Σ H	8.47	10.7	10.1	5.74	9.31	5.39	6.76
Na ₂ O	4.29	4.19	4.57	4.56	4.37	4.33	4.91	(La/Yb) _N	11.5	12.5	10.4	4.97	8.58	6.61	8.68
K ₂ O	2.43	2.45	2.51	1.94	2.06	2.37	2.40	(La/Sm) _N	4.47	4.06	3.99	2.86	5.45	2.40	2.57
P ₂ O ₅	0.17	0.16	0.16	0.15	0.18	0.19	0.16	(Gd/Yb) _N	2.23	2.11	1.80	1.19	1.05	2.04	2.30
烧失量	2.66	2.54	2.06	2.27	2.57	2.14	2.01	δ Eu	0.98	0.94	0.94	0.91	1.13	0.72	0.80
总量	100.8	100.7	99.7	99.6	100.7	99.8	99.7	Rb	55.5	71.3	58.6	63.6	48.0	55.9	63.5
K ₂ O+Na ₂ O	6.90	6.80	7.23	6.66	6.60	6.85	7.47	Sr	507	270	517	339	310	634	549
K ₂ O/Na ₂ O	0.57	0.59	0.55	0.43	0.47	0.55	0.49	Ba	472	537	571	522	554	1364	504
Fe ₂ O ₃ +FeO	4.63	4.79	3.71	3.46	4.24	4.17	3.13	Th	4.83	4.45	4.80	6.83	7.48	5.33	5.06
A/CNK	1.06	1.09	1.05	0.98	1.09	1.04	0.95	U	0.97	0.95	1.20	0.75	1.80	0.70	1.40
刚玉(C)	1.22	1.74	1.07	0.04	1.74	0.93	0	Nb	8.86	5.39	5.40	4.69	6.06	4.19	4.87
t_{Zr} (℃)	814	809	789	815	820	807	812	Ta	0.60	0.40	0.40	0.29	0.49	0.24	0.28
La	19.8	17.6	19	14.7	21.2	13.9	14.9	Zr	154	143	119	154	157	142	166
Ce	37.9	40.2	45.1	27.7	36.8	26.9	30.5	Hf	4.70	4.30	4.00	5.99	5.28	4.54	4.26
Pr	4.58	4.03	4.45	3.58	4.00	3.53	3.83	V	49.4	59.6	47.8	58.0	55.6	51.0	54.0
Nd	15.2	12.83	14.36	15.3	14.1	15.7	16.7	Ga	18.1	17.9	16.0	18.1	13.6	12.6	13.1
Sm	2.79	2.73	3.00	3.24	2.45	3.65	3.65	Y	13.8	8.46	10.78	17.2	13.7	25.0	21.6
Eu	0.96	0.81	0.90	0.94	0.87	0.86	0.93	Cr	64.7	50.9	21.8	48.9	46.3	52.5	49.5
Gd	3.19	2.47	2.75	2.94	2.16	3.58	3.29	Co	15.3	11.7	10.8	16.6	15.1	18.3	14.9
Tb	0.47	0.39	0.44	0.56	0.38	0.72	0.63	Ni	53.6	29.0	14.8	18.23	14.4	18.64	17.6
Dy	2.63	1.95	2.29	3.17	2.14	4.21	3.63	Th/Ta	8.05	11.1	12.0	23.3	15.4	22.2	17.8
Ho	0.53	0.39	0.46	0.65	0.45	0.46	0.45	Ta/Yb	0.52	0.42	0.32	0.15	0.29	0.17	0.24

注: $t_{Zr}/^{\circ}\text{C} = T_{Zr}/\text{K} - 273.15 = \frac{12900}{2.95 + 0.85M + \ln \frac{496000}{\text{Zr}_{\text{melt}}/10^{-6}}} - 273.15$; M 为全岩 $\frac{n(\text{Na}) + n(\text{K}) + 2n(\text{Ca})}{n(\text{Al}) \cdot n(\text{Si})}$, 计算中令 $n(\text{Si}) + n(\text{Al}) + n(\text{Fe}) + n(\text{Mg}) + n(\text{Ca}) + n(\text{Na}) + n(\text{K}) + n(\text{P}) = 1$; Zr_{melt} 为熔体中 Zr 含量 (Watson et al., 1983; 参见熊双才等, 2019; 张征峰等, 2011)。

特别是活动性强的元素的影响, 排除受影响较大的元素。通常被认为活动性差的元素(Si、Ti、Fe), 样品随蚀变程度升高它们基本能保持含量稳定, Cr、Ni 均与 LOI 之间没有明显相关性, 表明这些元素受热液活动的影响比较小; 通常认为碱金属、碱土金属元素(Na、K、Al、Mg)和大离子亲石元素(Rb、Sr、Ba、Pb、U 等)具有极高的活动性, 而高场强元素(Nb、Ta、Zr、Hf、REE 等)活动性很低 (Humphris and Thompson, 1978), 样品除大离子亲石元素存在一个异常点表现出了一定的活动性(经查证为 NWP304 样品 Ba 含量高达 1364×10^{-6} 所导致)外, 其它整体

保持稳定, 说明样品基本没有受热液蚀变的影响, 能够代表岩石的地球化学特征。

5.2.2 主量元素

南窝铺铀矿床英安岩 $\text{SiO}_2 = 65.49\% \sim 65.82\%$, $\text{K}_2\text{O} = 1.94\% \sim 2.51\%$, $\text{Na}_2\text{O} = 4.19\% \sim 4.91\%$, $(\text{K}_2\text{O} + \text{Na}_2\text{O}) = 6.60\% \sim 7.47\%$, $\text{K}_2\text{O}/\text{Na}_2\text{O} = 0.43 \sim 0.59$, 在 TAS 图解(图 8a)上落入亚碱性系列英安岩范围内, 在 $\text{SiO}_2-\text{K}_2\text{O}$ (图 8b)上落入钙碱性系列范围内; $(\text{FeO}+\text{Fe}_2\text{O}_3) = 3.13\% \sim 4.79\%$, $\text{MgO} = 1.25\% \sim 1.79\%$, $\text{TiO}_2 = 0.49\% \sim 0.67\%$, 在 AFM 图解(图 9a)上落入钙碱性系列范围内; $\text{Al}_2\text{O}_3 = 15.74\% \sim$

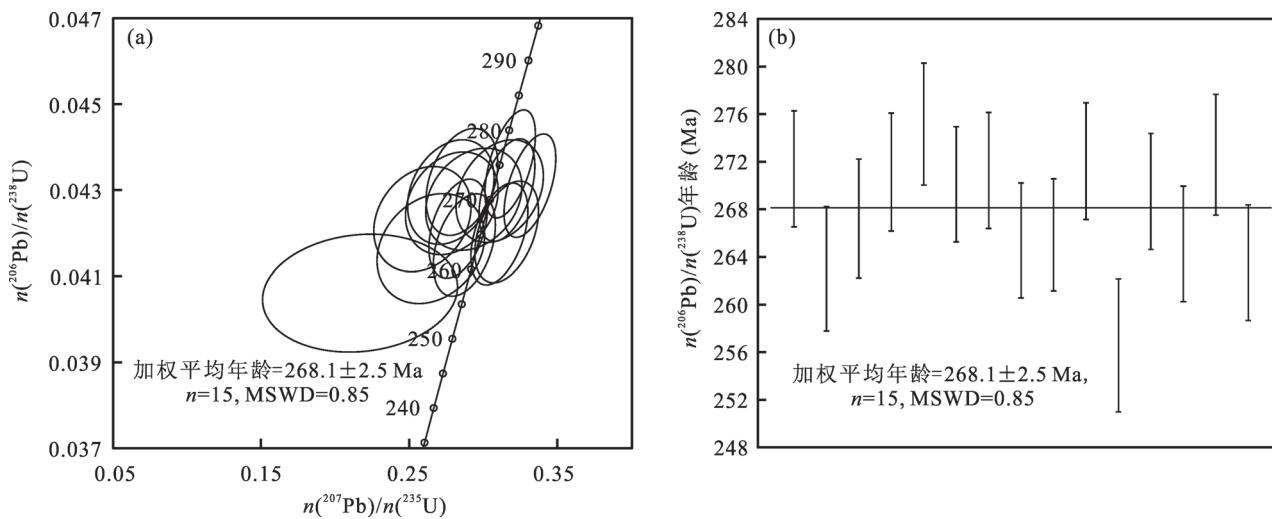


图 6 内蒙古赤峰南窝铺铀矿床英安岩锆石 U-Pb 年龄谐和图

Fig. 6 U-Pb age concordia diagram of dacite from the Nanwopu uranium deposit, Chifeng area, Inner Mongolia

16.77%, CaO = 2.87% ~ 3.59%, A/CNK = 0.95 ~ 1.09(平均 1.04, <1.1), 标准矿物刚玉(C)含量为 0 ~ 1.74%, 平均值为 0.96%, 具有 I 型英安岩的特征; MnO 含量为 0.04% ~ 0.05%, P₂O₅ 含量为 0.15% ~ 0.19%。

5.2.3 微量元素

南窝铺铀矿床英安岩稀土元素总量较低, $\Sigma \text{REE} = 76.5 \times 10^{-6} \sim 95.4 \times 10^{-6}$ (平均 85.9×10^{-6}), 配分模式呈轻稀土富集右倾型, 轻重稀土分馏明显 ($\text{La/Yb}_{\text{N}} = 4.97 \sim 12.52$ (平均 8.95)), 存在微弱的 Eu 异常 $\delta \text{Eu} = 0.72 \sim 1.13$ (平均 0.94)。球粒陨石标准化的稀土元素配分模式图(图 10a)上显示铕不亏损的右倾轻稀土富集特征, 与安第斯型钙碱性系列基本一致。原始地幔标准化蛛网图上, 英安岩富集大离子亲石元素(LILE) Rb、Ba、K、Sr 和强烈亏损高场强元素(HFSE) Nb、Ta、Ti(图 10b), 原始地幔标准化配分模式与安第斯型钙碱性系列岩石特点基本一致。

6 地质意义

6.1 地质时代

本次研究的南窝铺铀矿床英安岩 SHRIMP 锆石 U-Pb 年龄为 $268.1 \pm 2.5 \text{ Ma}$, 结合安山质角砾凝灰岩 SHRIMP 锆石 U-Pb 年龄为 $277.1 \pm 9 \text{ Ma}$ (纪宏伟, 2015), 指示南窝铺铀矿床赋矿安山岩—英安岩—流纹岩组合形成于 $268 \sim 277 \text{ Ma}$, 与红山子—广兴铀成矿亚带东部盔甲山一带额里图组角闪安山岩的

LA-ICP-MS 锆石 U-Pb 年龄($274 \pm 2 \text{ Ma}$; 彭斌等, 2016)一致, 为将赋矿安山岩—英安岩—流纹岩组合归入额里图组提供了年代学证据。根据 2018/08 版《国际年代地层表》(樊隽轩等, 2018), 上二叠统(乐平统)—中二叠统(瓜德鲁普统)、中二叠统(瓜德鲁普统)—下二叠统(乌拉尔统)和二叠系—石炭系的界线分别定为 $259.1 \pm 0.5 \text{ Ma}$ 、 $272.95 \pm 0.11 \text{ Ma}$ 和 $298.9 \pm 0.15 \text{ Ma}$, 指示南窝铺铀矿床赋矿安山岩—英安岩—流纹岩组合的地质时代为早二叠世—中二叠世早期。

南窝铺铀矿床赋矿英安岩的年龄确定, 不仅为红山子—广兴铀成矿亚带晚侏罗世新民组流纹岩—碱性流纹岩组合之下还存在早二叠世—中二叠世早期额里图组赋矿火山岩系提供了年代学证据, 而且该赋矿层位明显低于中国东部与火山岩有关的热液型的赋矿层位(巫建华等, 2017b), 是中国东部与火山岩有关的热液型铀矿新的赋矿层位。

6.2 岩石类型

南窝铺铀矿床英安岩的 SiO₂ 含量为 64.16% ~ 66.66%, Na₂O > 3.2%, K₂O/Na₂O 值小于 1, 铝饱和指数 A/CNK $\left[\frac{n(\text{Al}_2\text{O}_3)}{n(\text{CaO})+n(\text{Na}_2\text{O})+n(\text{K}_2\text{O})} \right]$ 小于 1.1, 说明岩石属于准铝质—弱过铝质 I 型火山岩。在 K₂O—Na₂O(图 9b)中, 样品投点也落在 I 型花岗岩区。南窝铺铀矿床英安岩的锆石饱和温度为 $789^\circ\text{C} \sim 827^\circ\text{C}$, 平均为 812°C , 明显低于红山子盆地 A 型流纹岩($951^\circ\text{C} \sim 988^\circ\text{C}$, 巫建华等, 2016)的锆石

饱和温度,也低于 A 型花岗岩平均温度 833℃(刘昌实等,2003),也接近高分异的 I 型花岗岩平均温度 781℃(King et al., 1997)。南窝铺铀矿床英安岩的 A/CNK 比值小于 1.1,属准铝质—弱过铝质花岗岩类(高栋等,2018), P_2O_5 含量很低且随着 SiO_2 含量增高而降低,呈负相关关系,具有 I 型花岗岩特征(Wolf et al., 1994; Chappell, 1999; 李子昊, 2018)。在 SiO_2 —Ce 和 Al_2O_3 —Ga 图解(图 11)中落入 I 型花

岗岩范围,明显不同于红山子—广兴铀成矿亚带(以红山子盆地为代表)新民组 A 型流纹岩,相比其他地区额里图组流纹岩,均为 I 型花岗岩(崔玉良等,2019)。可见,南窝铺铀矿床额里图组英安岩具有 I 型花岗岩特征,为有别于新民组 A 型流纹岩一类新的赋矿围岩类型。

6.3 构造环境

在微量元素 Yb—Ta、Y—Nb、Rb—Y + Nb 和

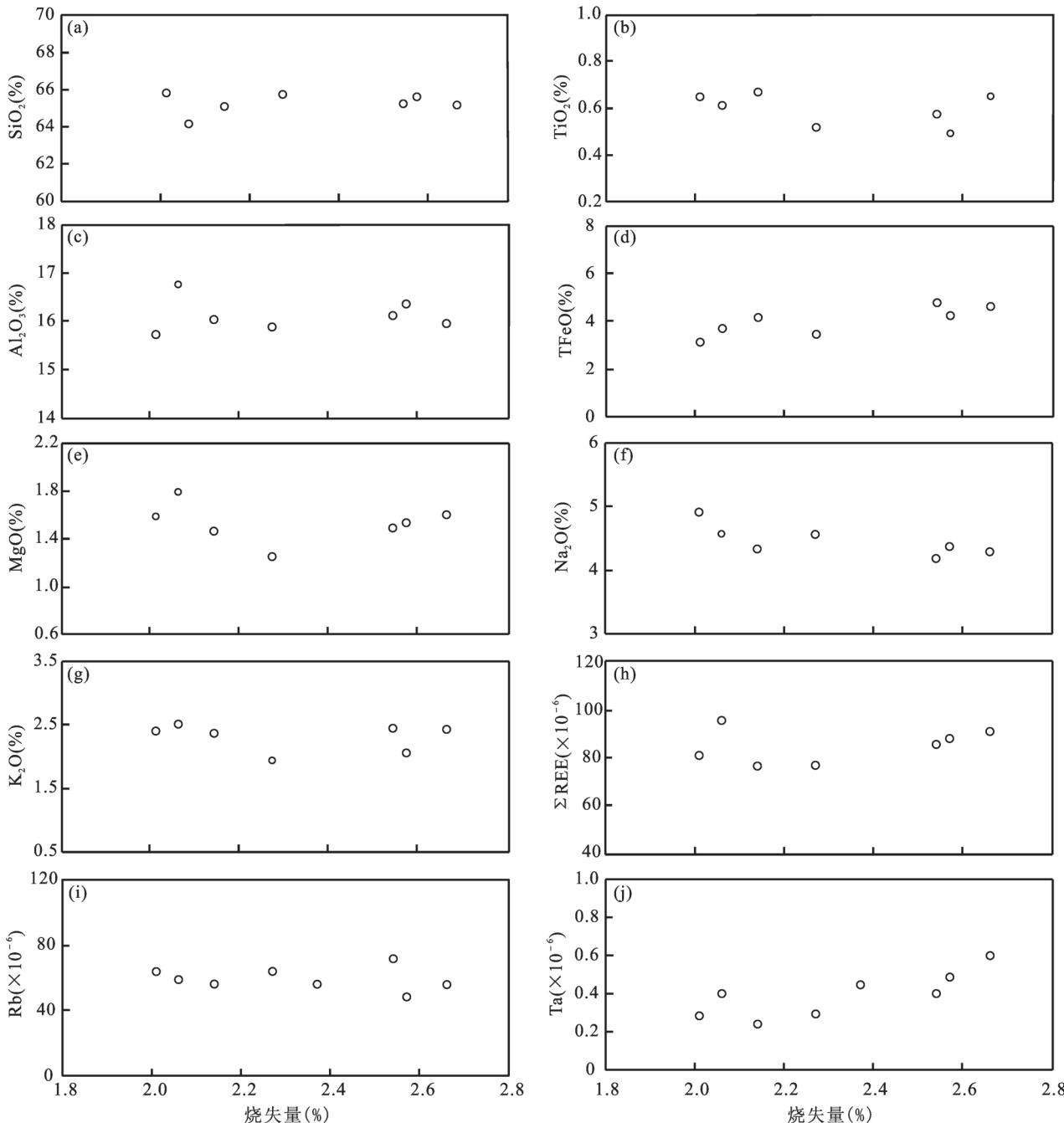


图 7 内蒙古赤峰南窝铺铀矿床英安岩烧失量—主微量元素协变图解

Fig. 7 Bivariate diagrams of LOI—elements of dacite from the Nanwopu uranium deposit, Chifeng area, Inner Mongolia

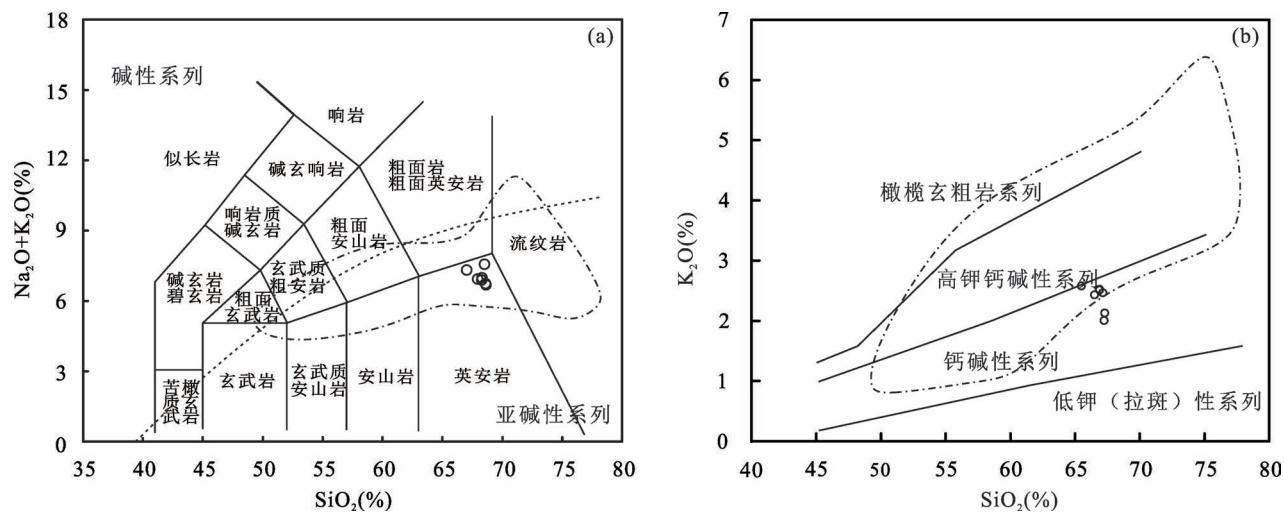


图 8 内蒙古赤峰南窝铺铀矿床英安岩岩石分类判别图解(虚线范围代表其他地区额里图组火山岩,数据来源:曹代勇等,2014;董晓杰等,2016;彭斌等,2016;崔玉良等,2019)。(a) TAS 图解(底图据 Middlemost, 1994);(b) SiO₂—K₂O 图解(底图据 Rollison, 1993)

Fig. 8 Rock classification and discrimination diagrams of dacite from the Nanwopu uranium deposit, Chifeng area, Inner Mongolia (The dashed area is the literature data of the Elitu Formation from other areas, data sources: Cao Daiyong et al. , 2014&; Dong Xiaojie et al. , 2016&; Peng Bin et al. , 2016&; Cui Yuliang et al. , 2019&). (a) TAS diagram(after Middlemost, 1994); (b) SiO₂—K₂O diagram(after Rollison, 1993)

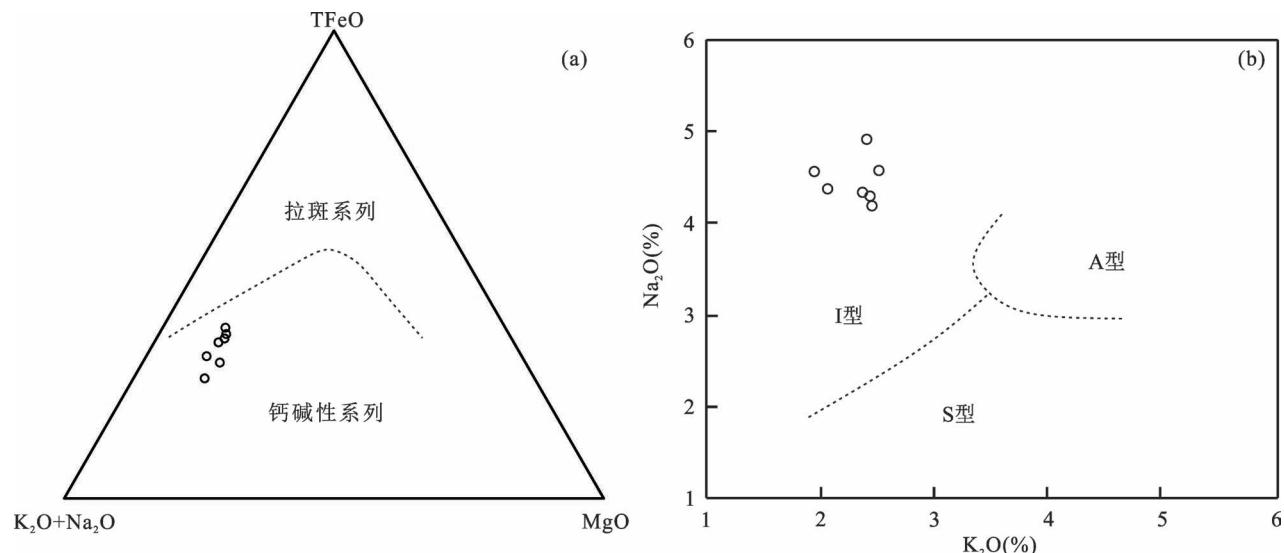


图 9 内蒙古赤峰南窝铺铀矿床英安岩 AFM 图解(a,底图据 Irvine et al. , 1971)和 Na₂O—K₂O 图解(b,底图据 Middlemost, 1986)

Fig. 9 Diagrams of AFM (a, after Irvine et al. , 1971) and Na₂O—K₂O (b, after Middlemost, 1986) of dacite from the Nanwopu uranium deposit, Chifeng area, Inner Mongolia

Rb—Yb+Ta 构造环境判别图解中(图 12), 南窝铺铀矿床的英安岩均投影于火山弧花岗岩区, 反映其形成于活动大陆边缘环境。在稀土元素球粒陨石标

准化配分模式图和微量元素原始地幔标准化配分模式图(图 10)上英安岩显示与安第斯型钙碱性系列基本一致的特征。因此, 南窝铺铀矿床的英安岩可

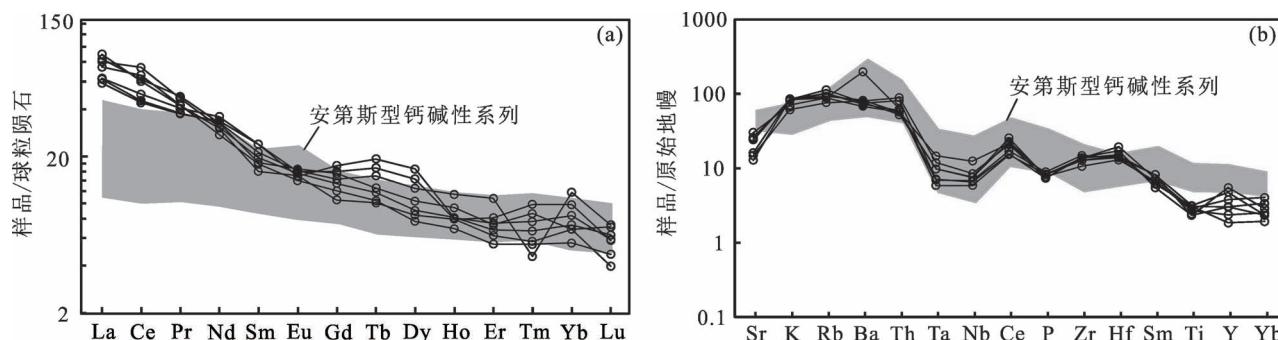


图 10 内蒙古赤峰南窝铺铀矿床英安岩球粒陨石标准化稀土元素配分模式图(a)和原始地幔标准化微量元素蛛网图(b)
(标准化值据 Sun and McDonough, 1989; 安第斯型钙碱性系列数值据 Gutiérrez, 2005)

Fig. 10 Chondrite-normalized REE distribution pattern (a) and Primitive mantle-normalized spidergram (b) of dacite from the Nanwopu uranium deposit (Normalizing values from Sun and McDonough, 1989; Andean-type calcium-alkaline series data from Gutiérrez, 2005)

能为安第斯型活动大陆边缘火山弧的产物,与华北古板块北缘集宁北部商都县西井子镇一带早—中二叠世(267~272 Ma)苏吉组酸性火山岩(董晓杰等,2016)、赤峰北部朝阳沟村早二叠世流纹岩(崔玉良等,2019)、赤峰铭山隆起带早二叠世额里图组安山岩(彭斌等,2016)、赤峰翁牛特旗早二叠世流纹岩—安山岩—流纹质—英安质凝灰岩组合(曹代勇等,2014)、吉林中部早二叠世大河深组流纹岩—英安岩—粗面英安岩组合和延吉地区早二叠世关门咀子组玄武安山岩形成的构造环境一致(曹花花等,2012; Yu Qian et al., 2014)。

南窝铺铀矿床赋矿英安岩是安第斯型活动大陆边缘火山弧的产物,不仅与红山子—广兴铀成矿亚带晚侏罗世新民组板内拉张环境下的流纹岩—碱性流纹岩组合(巫建华等,2013,2016,2017a,2017b; 彭啟辉,2015; 解开瑞等,2016; 黎伟等,2017; 姜山等,2018)、沽源—丰宁铀成矿亚带早白垩世张家口组板内拉张环境下的流纹岩—粗面岩组合(巫建华等,2014,2015,2017c; 张雅菲等,2016; 夏应冰等,2016; 林天发等,2019)不同,而且与中国东部赣杭铀成矿带早白垩世武夷群板内拉张环境下的流纹岩—粗面岩组合(刘飞宇等,2009; 巫建华等,2011,

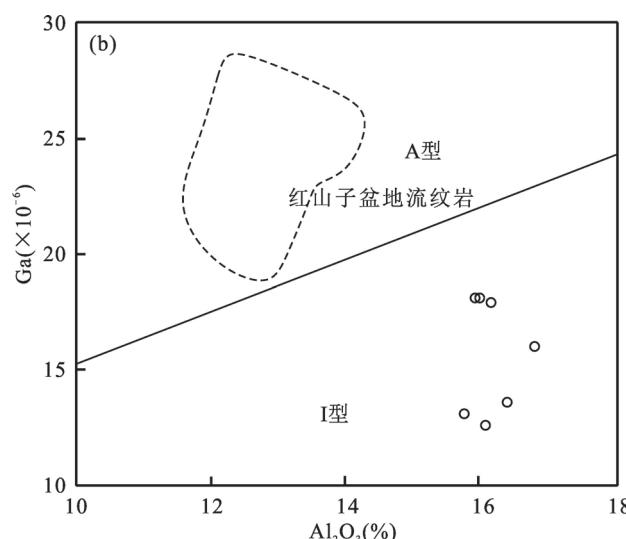
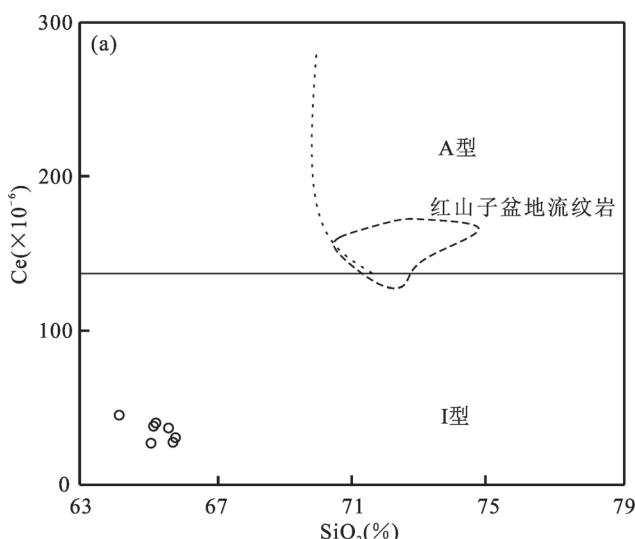


图 11 内蒙古赤峰南窝铺铀矿床英安岩 SiO_2 —Ce 图(a) 和 Al_2O_3 —Ga 图(b)
(红山子盆地流纹岩数据来源:巫建华等,2016)

Fig. 11 SiO_2 —Ce (a) and Al_2O_3 —Ga (b) diagrams of dacite from the Nanwopu uranium deposit, Chifeng area, Inner Mongolia (Data of the Hongshanzi basin rhyolite are from Wu Jianhua et al., 2016&)

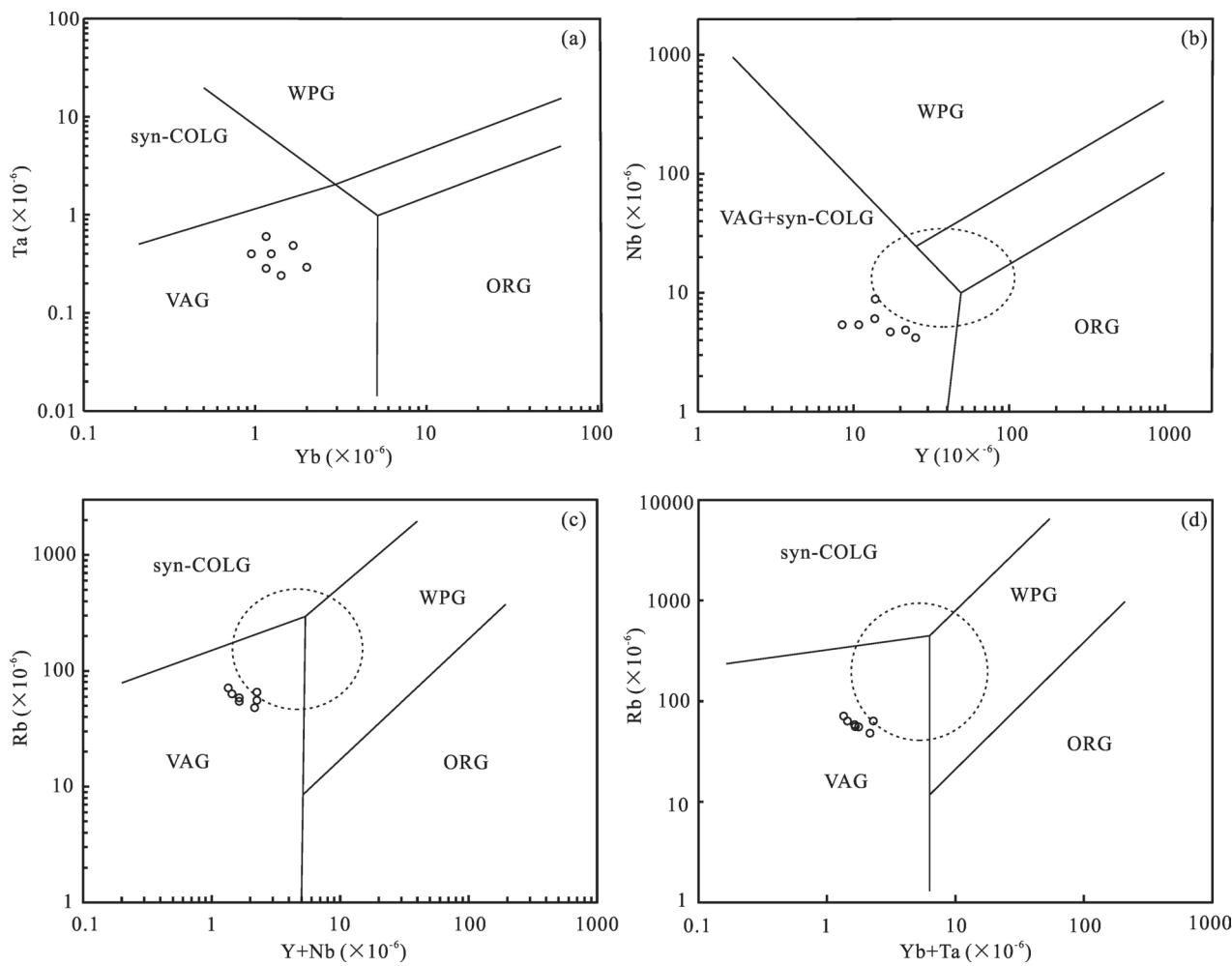


图 12 内蒙古赤峰南窝铺铀矿床英安岩构造判别图解

(底图据 Pearce et al., 1984; Roser et al., 1986; Murray, 1994; Gorton et al., 2000)

Fig. 12 Discrimination diagrams of tectonic environment for dacite from the Nanwopu uranium deposit, Chifeng area, Inner Mongolia (after Pearce et al., 1984; Roser et al., 1986; Murray, 1994; Gorton et al., 2000)

2017b; 韦昌袭, 2019)、武夷山铀成矿带早侏罗世余田群板内拉张环境下的流纹岩—玄武岩组合(冀春雨等, 2011; 项媛馨等, 2012)、晚白垩世兴宁群大量流纹岩—少量玄武岩组合(巫建华等, 2017b; 劳玉军等, 2018)不同, 是中国东部与火山岩有关的热液型铀矿一类新的构造环境下的赋矿火山岩组合。

7 结论

(1) 南窝铺铀矿床赋矿围岩英安岩 SHRIMP 镓石 U-Pb 年龄为 268.1 ± 2.5 Ma, 结合前人安山质角砾凝灰岩 277.1 ± 0.9 Ma 的镓石 U-Pb 年龄, 限定额里图组安山岩—英安岩—流纹岩组合的时代为早二叠世, 与中国东部中生代与火山岩有关的热液型铀矿赋矿围岩地质时代不同, 为新的赋矿层位;

(2) 南窝铺铀矿床额里图组英安岩为钙碱性系列, 属于 I 型火山岩, 具有明显不同于红山子—广兴铀成矿亚带下侏罗统新民组流纹岩的地球化学特征, 其明显富集 Rb、Ba、K、Sr 等大离子亲石元素, 亏损 Nb、Ta、Ti 等高场强元素;

(3) 南窝铺铀矿床额里图组英安岩为活动大陆边缘火山弧的产物, 相比已知的中国东部与火山岩有关的热液型铀矿床中生代赋矿围岩, 其构造环境明显不同。

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SHRIMP zircon U-Pb ages, geochemical characteristics and tectonic significance of dacite from the Nanwopu Uranium Deposit, Chifeng, Inner Mongolia

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Objectives: The Nanwopu uranium deposit is located in the northeastern section of the Guyuan—Hongshanzi uranium metallogenic belt. Dacite, which belongs to Elitu Formation, is the direct surrounding rock of the Nanwopu uranium deposit. The aim of this paper is to investigate the petrogenesis and tectonic significance of this rock suite.

Methods: Detailed geological survey, SHRIMP zircon U-Pb dating and whole-rock major and trace elements were carried out on the Nanwopu dacite in this paper.

Results: The SHRIMP zircon U-Pb dating of the Nanwopu dacite is 268.1 ± 2.5 Ma, it can be ascertained that the dacites are the products of Early Permian magmatic activities. The whole-rock geochemical analyses show that the Nanwopu dacite lie in middle K calc-alkaline fields, the Aluminium saturation index show the metaluminous and weak peraluminous features, integrated with enrichments of LREE and LILE (Rb, Ba, Th, K), depletions of Nb, Ta, P and Ti and Eu negative anomaly, belongs to I-type granites, suggesting plate subduction-related geochemical affinities.

Conclusions: Combined with the regional studies, the above geochemical signatures indicate that the Nanwopu dacite was formed in the volcanic arc tectonic environment on the active continental margin of the Hercynian tectonic stage (north margin of North China Craton). This is inconsistent with the tectonic environment drew from the other Mesozoic uranium deposits related to volcanic rocks in eastern China. The Lower Permian Erlitu Formation is favorable for the occurrence of uranium mineralization in eastern China.

Keywords: dacite; Lower—Middle Permian; Elitu Formation; uranium mineralization; geological significance

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